

SPARC DA Workshop

Derivation of wind fields from stratospheric trace gas observations

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MIPAS ESA team



Outline

- ARISE project
- Approach
- Skill analysis
- Species analysis
- Kiruna comparisons



ARISE project

- EU FP7/H2020 Collaborative Infrastructure Design Study
- Phase 1: 2012-2014; phase 2: 2015-2018
- Lead: CEA
- Middle Atmosphere: Linking models and observations

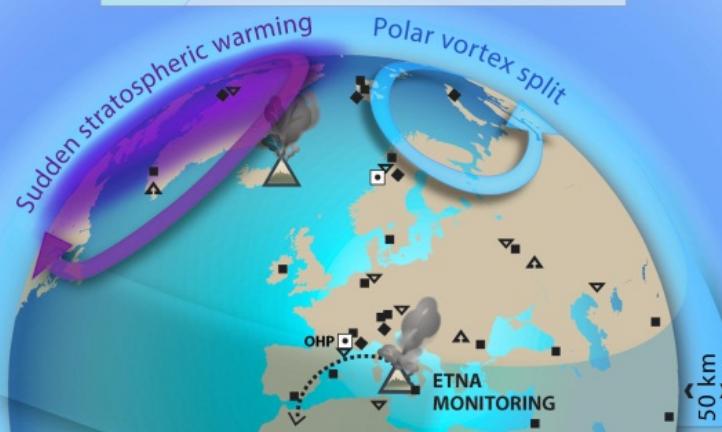
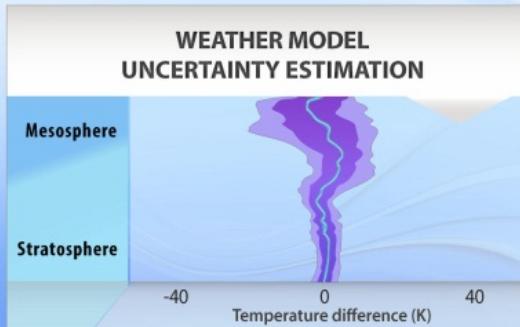
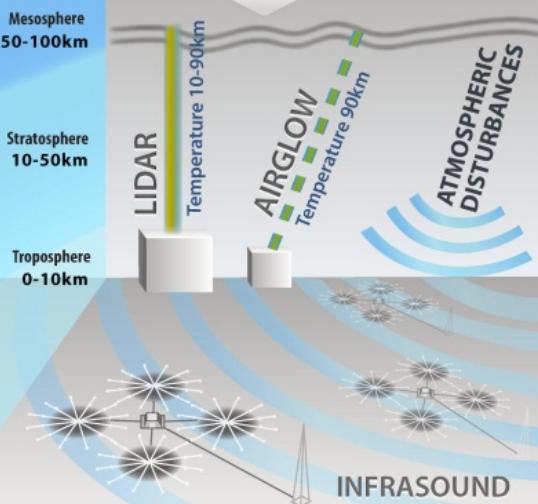
ARISE consortium : Partner's and member's logos



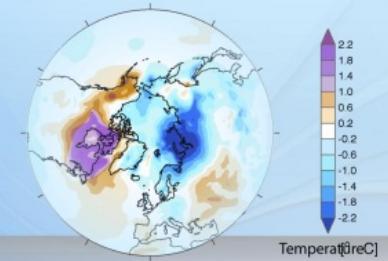
ARISE middle atmospheric dynamics - linking observations and models

ARISE MEASUREMENT STATION OBSERVATIONS ON MULTIPLE SCALES



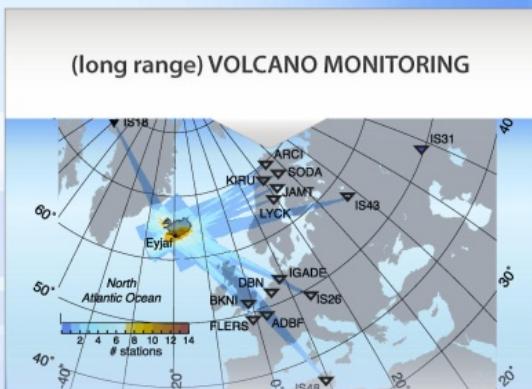
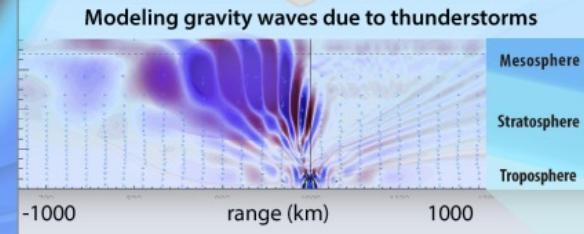
ARISE MODELING HIGH PERFORMANCE COMPUTING

TOWARDS WEATHER FORECAST IMPROVEMENT



Weather change 20 days following
a Sudden Stratospheric Warming

HIGH RESOLUTION ATMOSPHERIC DYNAMICS SIMULATION



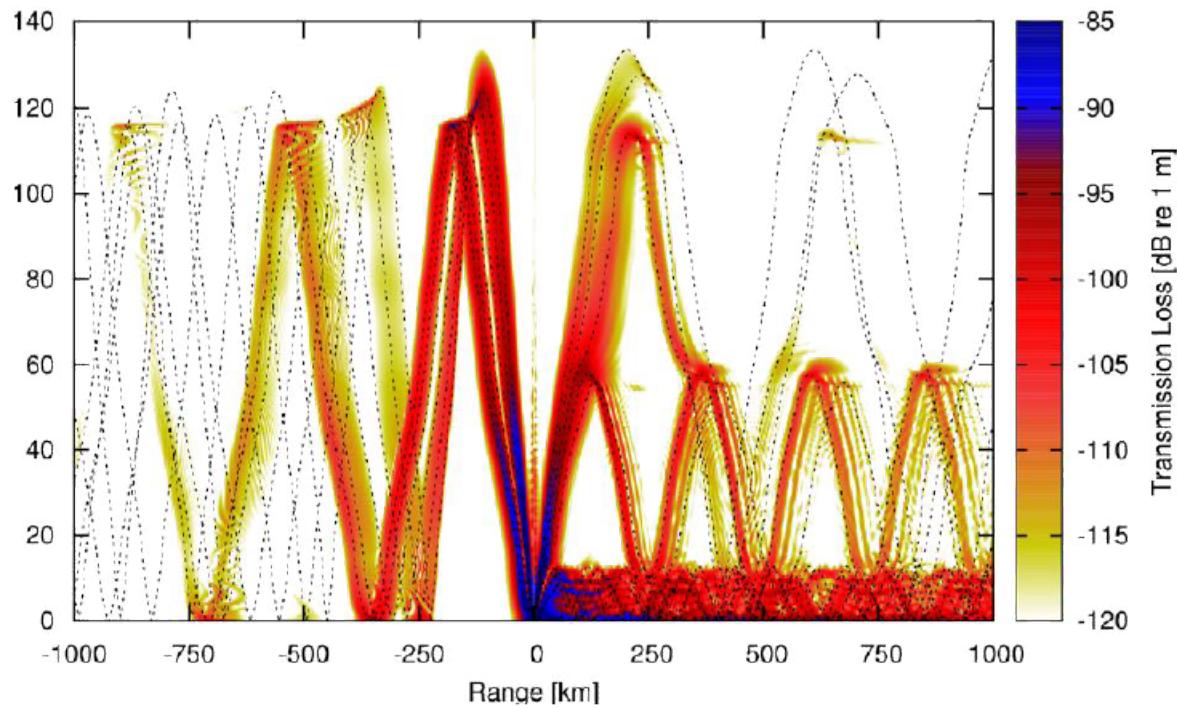
AVAILABLE TO PUBLIC



arise-project.eu

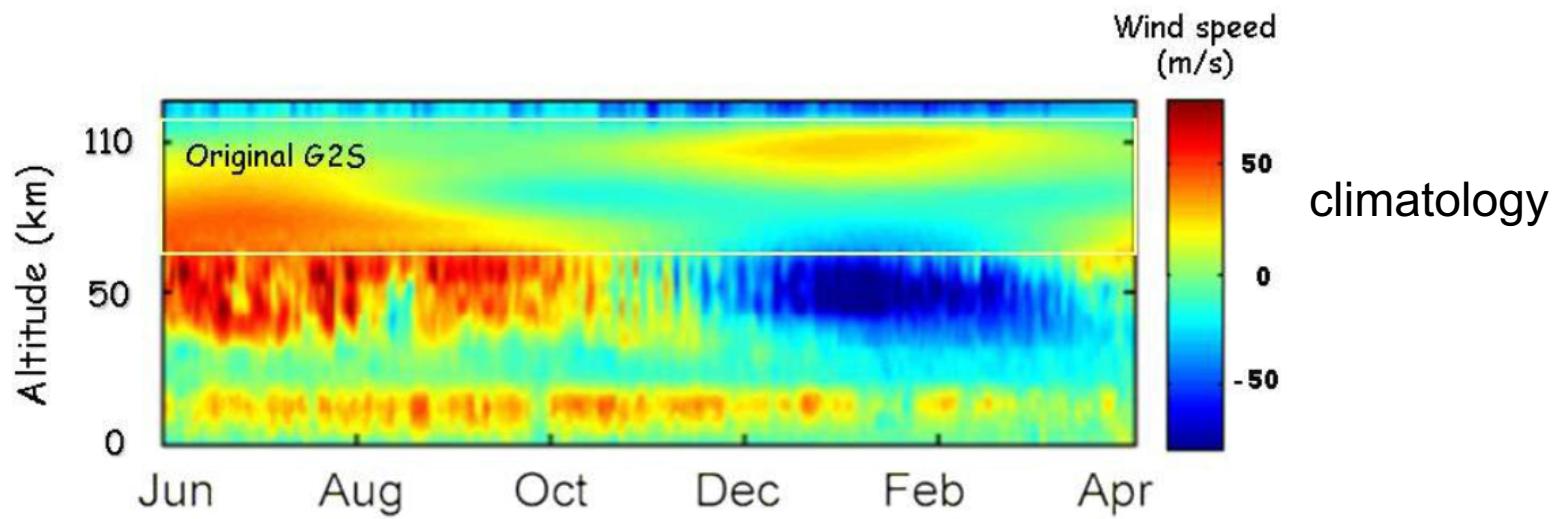
ARISE inversion methods: infrasound propagation

@Le Pichon et al., 2014

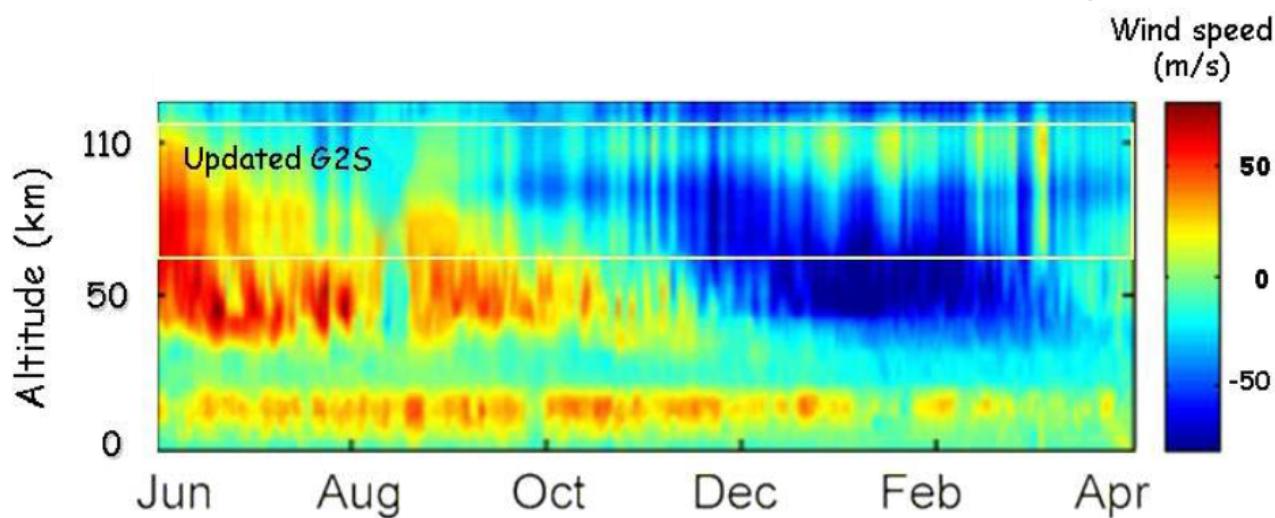


Typical IS propagation during boreal winter with wave guides near the Termopause, Stratopause and Tropopause

ARISE: reconstruction of upper stratospheric winds



@Le Pichon et al., 2014



Derivation of winds from trace gas observations

- Daley, 1995: 1-dimensional Kalman-Filter tracer experiments
- Grandpré, 2007: 4Dvar assimilation of chemical constituents
- Seman et al., 2009: O₃ profile assimilation with Méteo-France 4Dvar NWP OP-suite



This study

- 2D isentropic model to study stratospheric winds from tracer inversion
- GME-SACADA coupled model (Elbern et al., 2010) with new adjoint for advection scheme: global study with MIPAS observations
- Comparisons to local wind soundings: Kiruna radiosonde station



What can be expected?

- 2D-isentropic (const. wind) tracer model experiments:
 - 3h-lagged GME winds for Feb 1st, 2003
 - run isentropic model on 475K level with MIPAS-SACADA O3 tracer
 - compare 24h wind errors w. and wo. lag (skill loss)
 - calculate recover of initial winds with 6h inversion (inv. score)

RMS (Root Mean Squared) wind error scoring

Wind distortion

$$du = u(\text{perfect}) - u(\text{distorted})$$

rms wind error

$$\text{rms} = \sqrt{\frac{1}{N} \sum (du_i^2 + dv_i^2)}$$

24h forecast skill

(GME' with 3h-lagged winds)

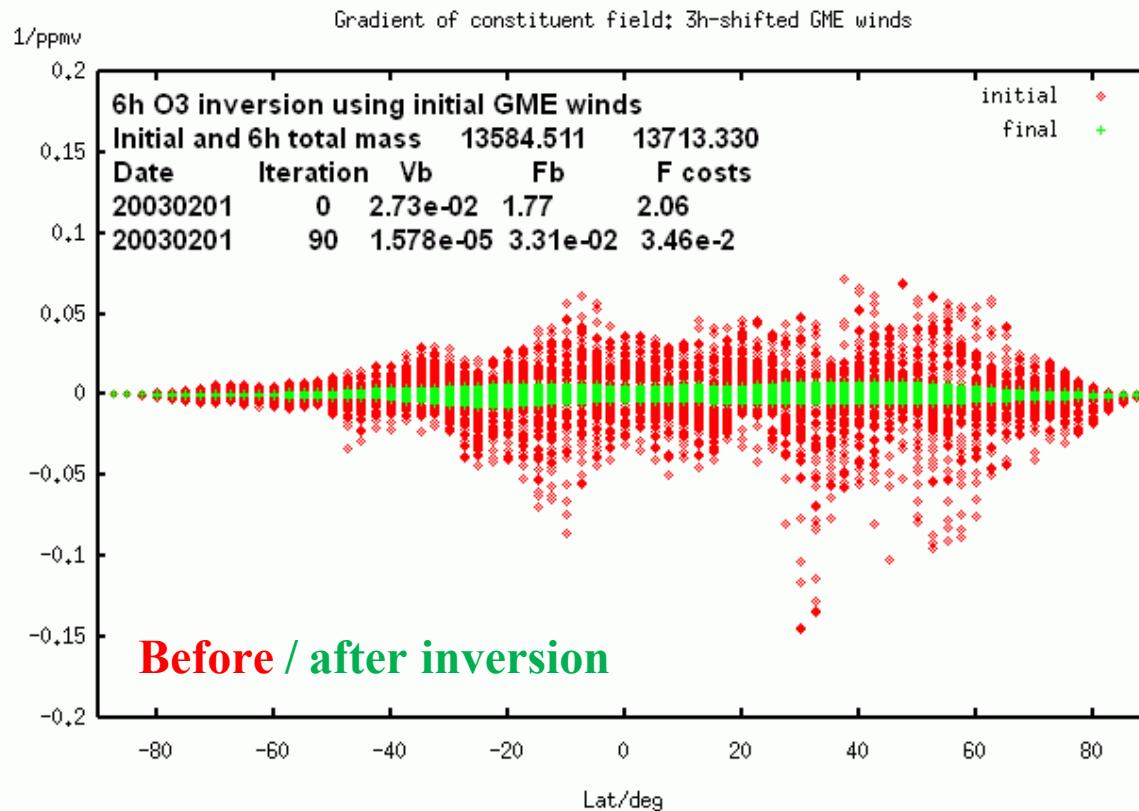
$$\text{fsk} = 1 - \frac{\text{rms}(GME' - ECMWF)}{\text{rms}(GME - ECMWF)}$$

6h inversion score

$$\text{isc} = 1 - \frac{\text{rms}(\text{initial winds after inversion})}{\text{rms}(\text{initial winds before inversion})}$$

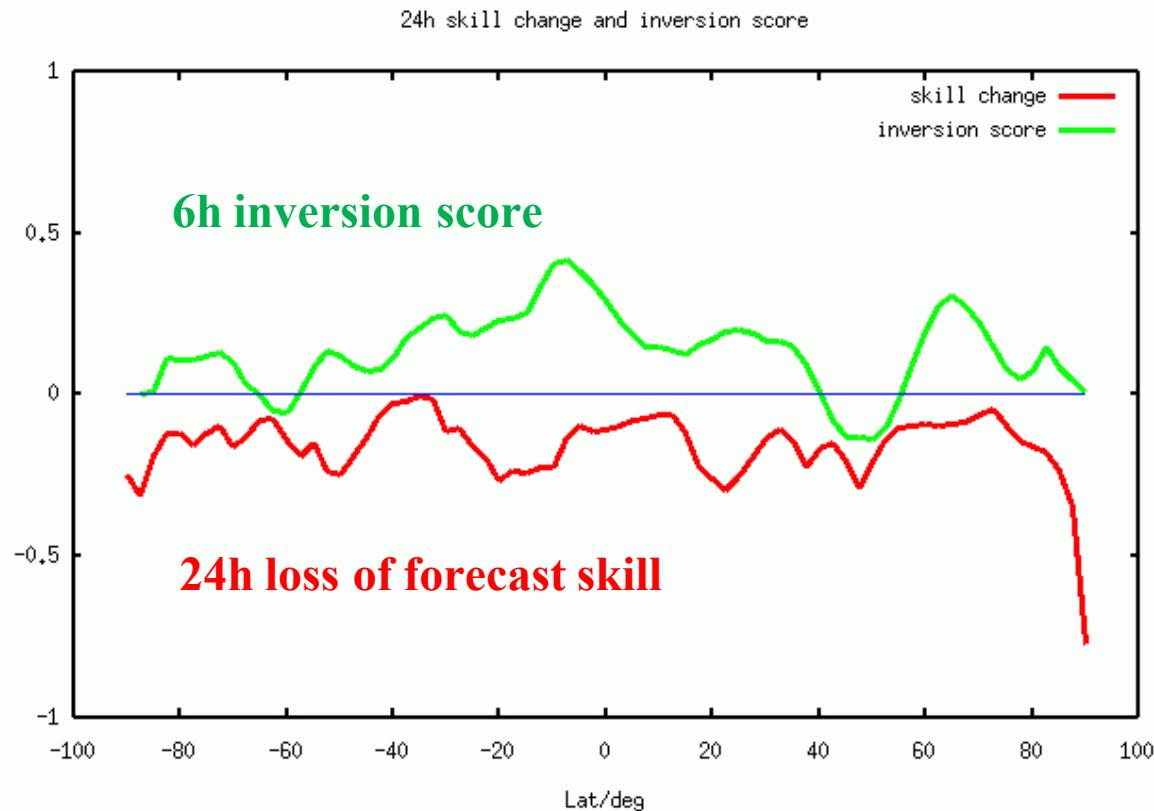


What can be expected?



Test shows that initial wind can be retrieved

What can be expected?



Summary isentropic 475K model

- Experiments with MIPAS O₃, CH₄ and N₂O for Feb. 2003 show
 - Inversion works best for low and high latitudes
 - Only weak differences between species
- Results comply with Seman, Grandpré et al.



GME-SACADA 4D-Var assimilation system

- GME
 - German global forecasting system 0-60 km altitude (42 levels)
 - icosahedral grid: ~250 km grid spacing
- SACADA
 - full stratospheric transport-chemistry scheme with adjoint
 - *New: adjoint of advection w.r.t. initial wind field*
- 4Dvar scheme
 - Weaver & Courtier incremental background-cov. modelling with
 - Horizontal and vertical lengthscales = 600 km, 3 km



Calculating wind-gradient of tracer cost function

$$\delta\mathbf{x} \equiv \mathbf{x} - \mathbf{x}_b ; \delta\mathbf{x} \equiv H_{adv}\delta\mathbf{u} ; \mathbf{d} \equiv \mathbf{y}^o - H\mathbf{x}_b ; \tilde{H} \equiv H \cdot H_{adv}$$
$$J(\delta\mathbf{u}) = \frac{1}{2} \delta\mathbf{u}^T B^{-1} \delta\mathbf{u} + \frac{1}{2} (\tilde{H}\delta\mathbf{u} - \mathbf{d})^T R^{-1} (\tilde{H}\delta\mathbf{u} - \mathbf{d})$$

$$\square := B^{-1/2} \delta\mathbf{u} \Rightarrow \delta\mathbf{u} = B^{1/2} \square \quad (\text{Weaver, Courtier 2001})$$

$$\nabla_{\delta\mathbf{u}} J = \nabla_{\delta\mathbf{u}} J_b + \nabla_{\delta\mathbf{u}} J_o = \square \cdot B^{T/2} \nabla_{\delta\mathbf{x}} J_o$$

$$\nabla_{\delta\mathbf{x}} J_o = \tilde{H}^T R^{-1} (\tilde{H}\delta\mathbf{u} - \mathbf{d}) = {H_{adv}}^T H^T R^{-1} (H\delta\mathbf{x} - \mathbf{d})$$

Initial gradient does not depend on background!



Advection operator and gradient of cost function

$$vmr_n = H_{adv}(vmr_{n-1}) = H_{hor} \circ H_{vert} (vmr_{n-1})$$

Semi-Lagrange first-order
| Diagnostic Euler-upward
|
 $\nabla_{\delta x} J_o = \tilde{H}^T R^{-1} (\tilde{H} \delta u - d) = {H_{adv}}^T H^T R^{-1} (H \delta x - d)$
|
adjoint



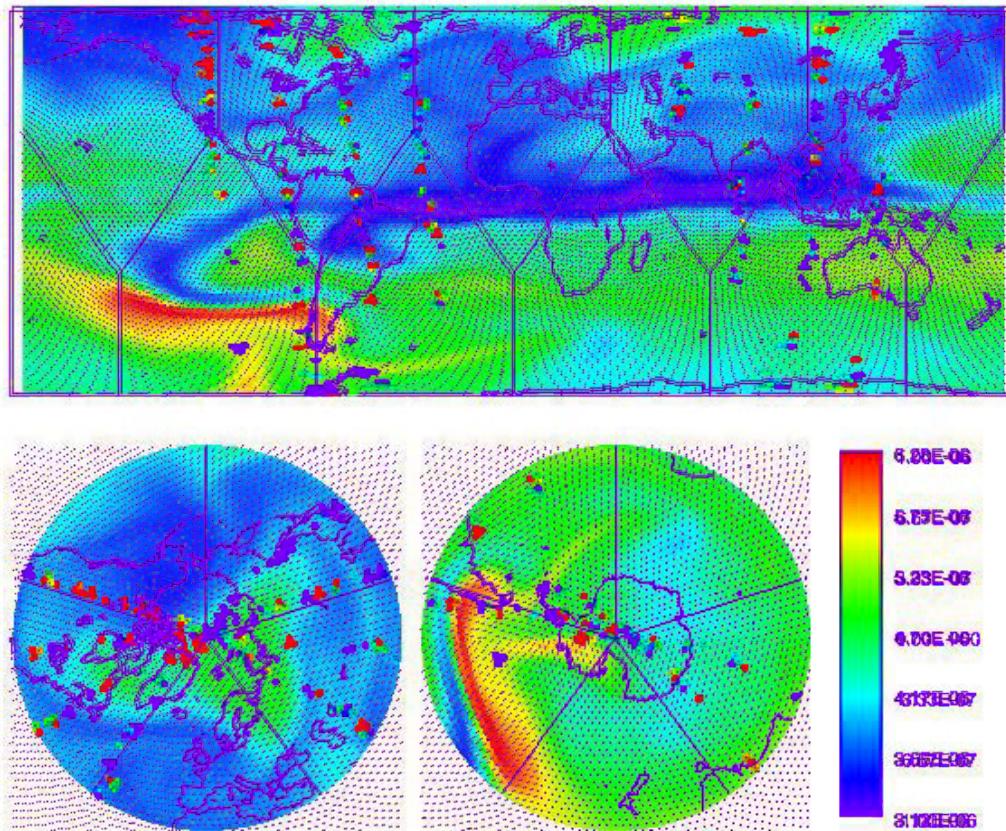
Inversion experiments Feb. 2003

- MIPAS (ESA): H₂O, HNO₃, N₂O, O₃, CH₄
- 6h assimilation window (max 15 iterations): 10% wind and vmr errors
- derive new initial wind field for next GME-SACADA fwd run cycle
- assess global cost function (FMOs) w. and wo. Inversion

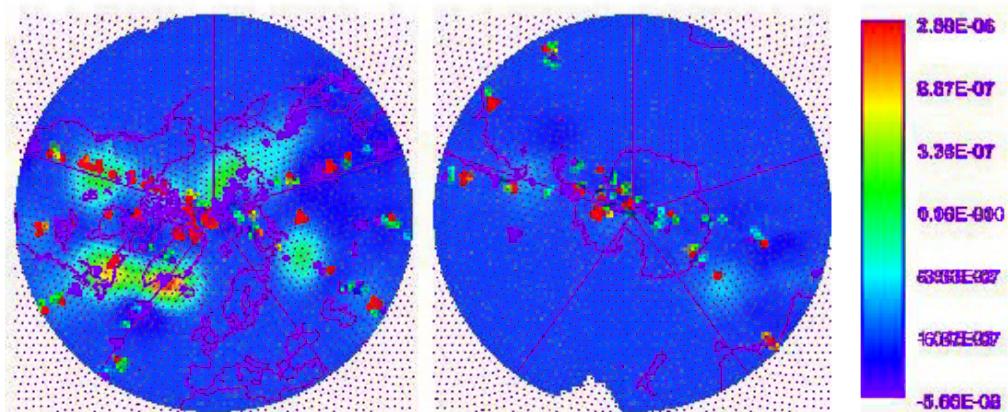
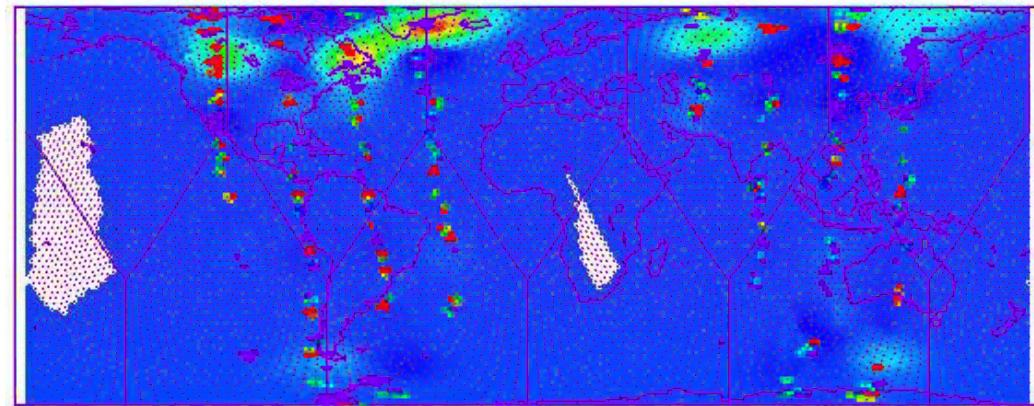
Meteo Februar northern stratosphere: SSW with vortex spilt in second half.



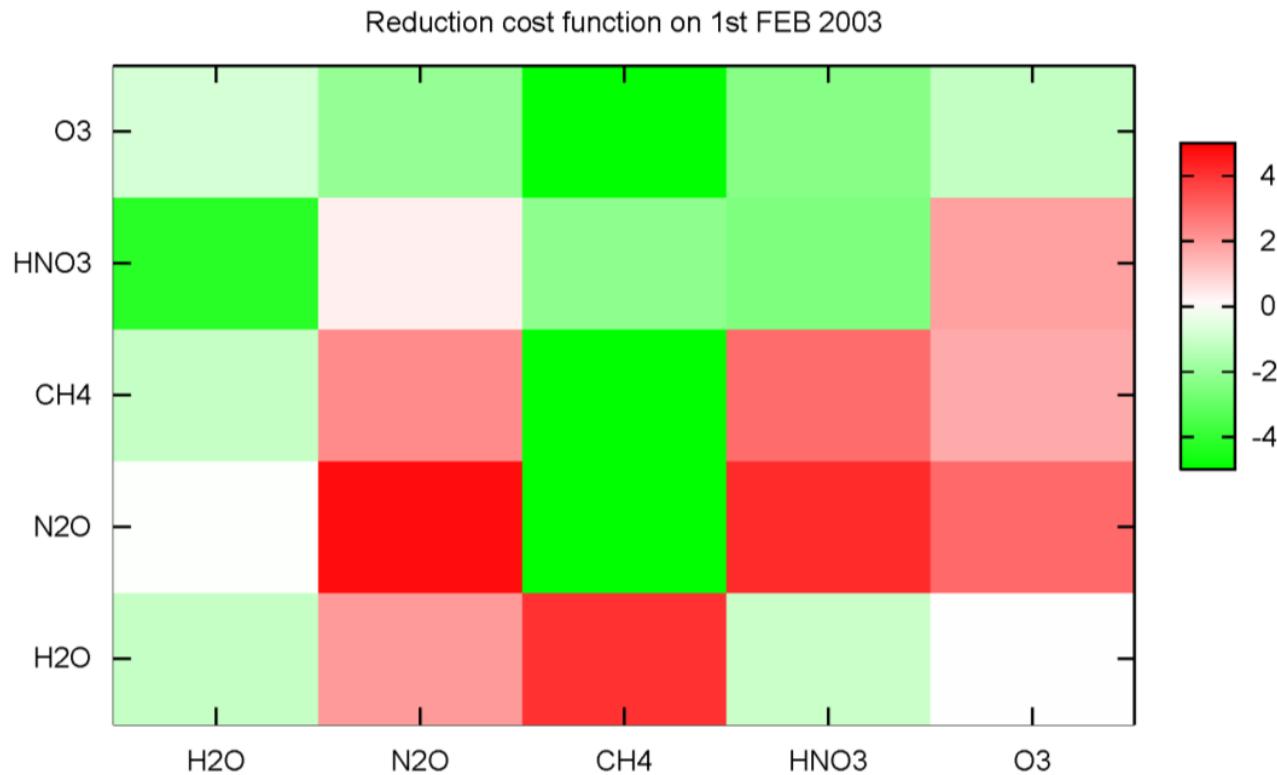
Sample 20 km H₂O field from MIPAS analysis



Correction of initial wind field

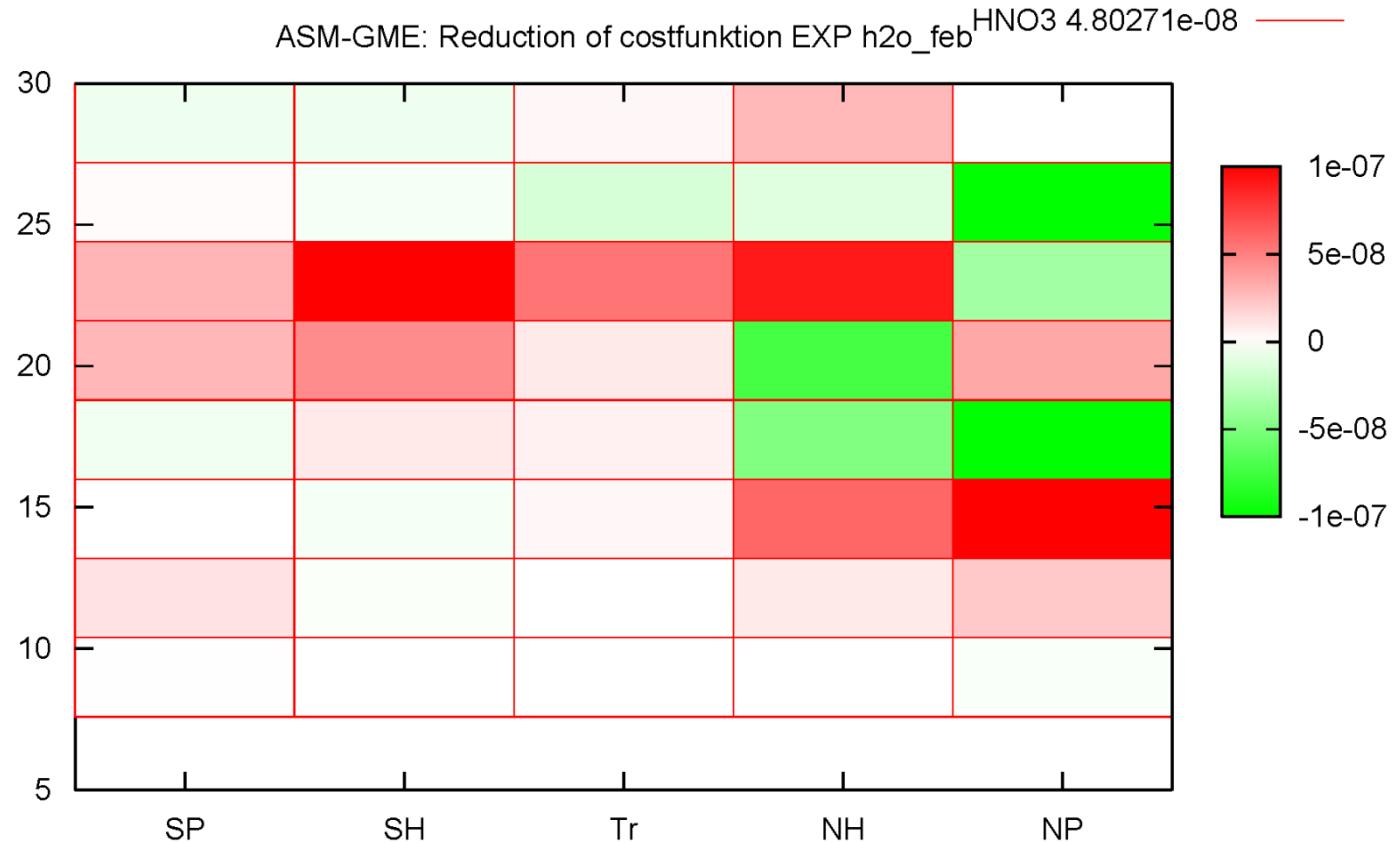


Sensitivity study for different MIPAS trace gases

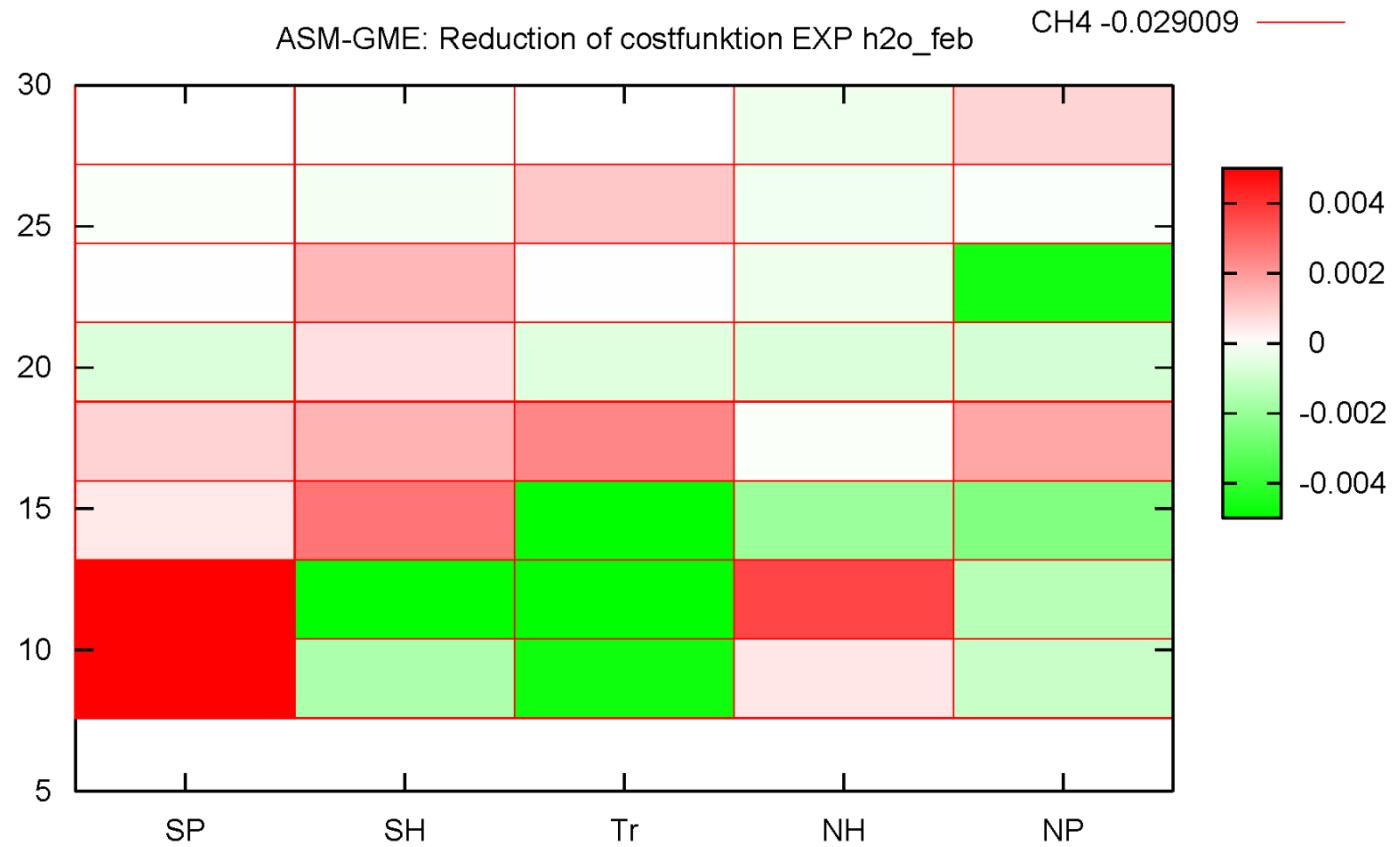


Reduction of cost function on 1st February, 2003: Y-axis shows assimilated species; X-axis shows species cost function. Note different scaling factors: O3= 2.e2; HNO3= 2.e7; CH4= 1.e4; N2O= 5.e4; H2O= 1.

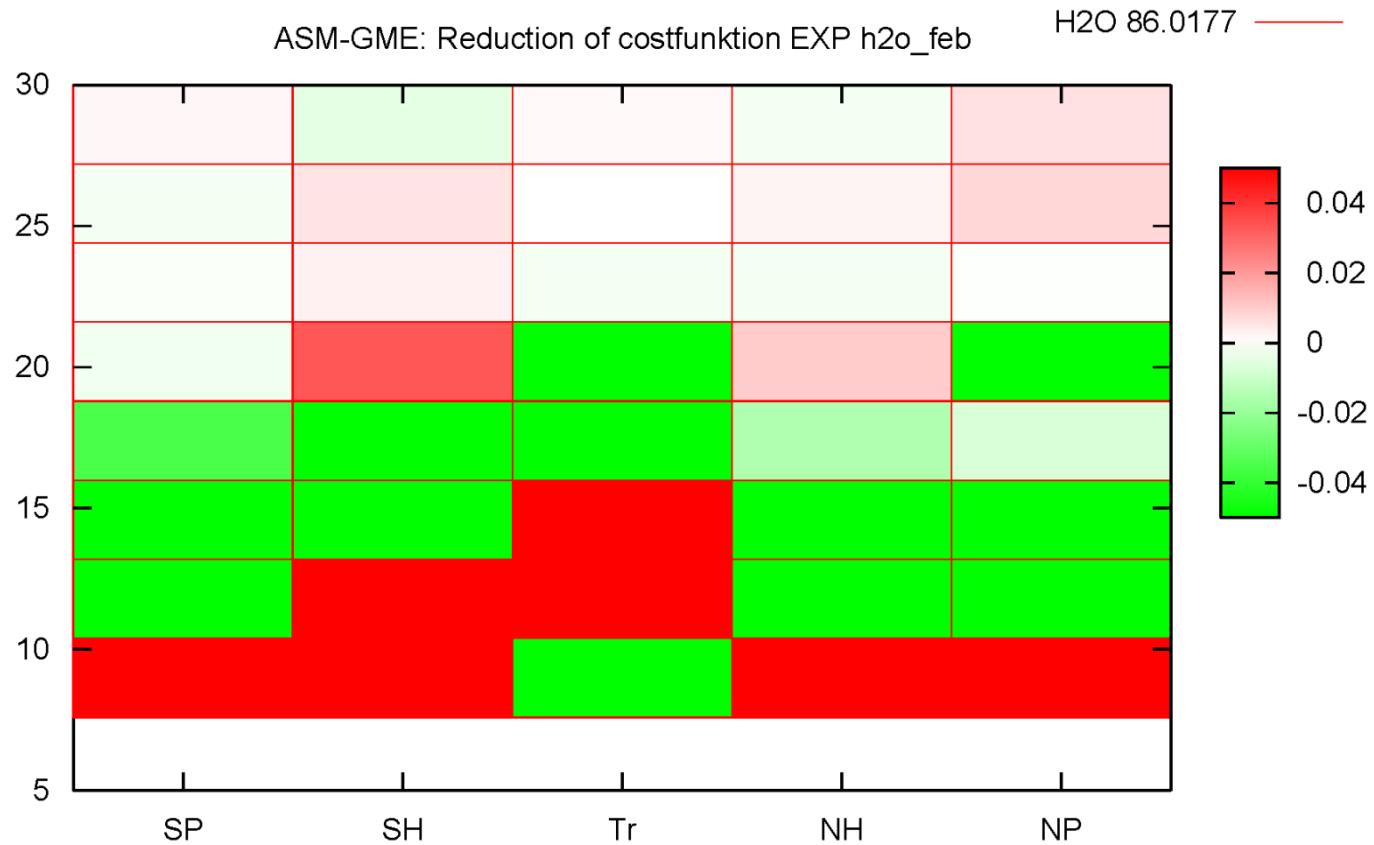
Reduction of cost function: HNO3 Feb. 2003



Reduction of cost function: CH4 Feb. 2003



Reduction of cost function: H2O Feb. 2003



Summary GME-SACADA experiments Feb. 2003

- results are not conclusive (cost function versus wind corrections)
- missing global convergence on some days
- tendency for vertical oscillations – why?
- covariance modelling too coarse? Error in inversion model?

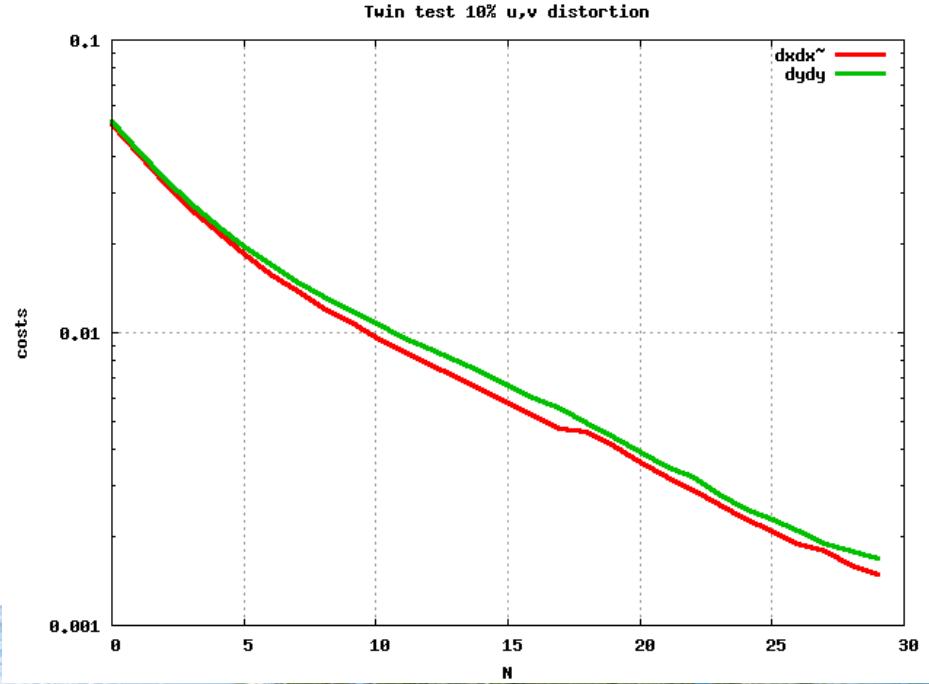


Offline assimilation experiments March 2003

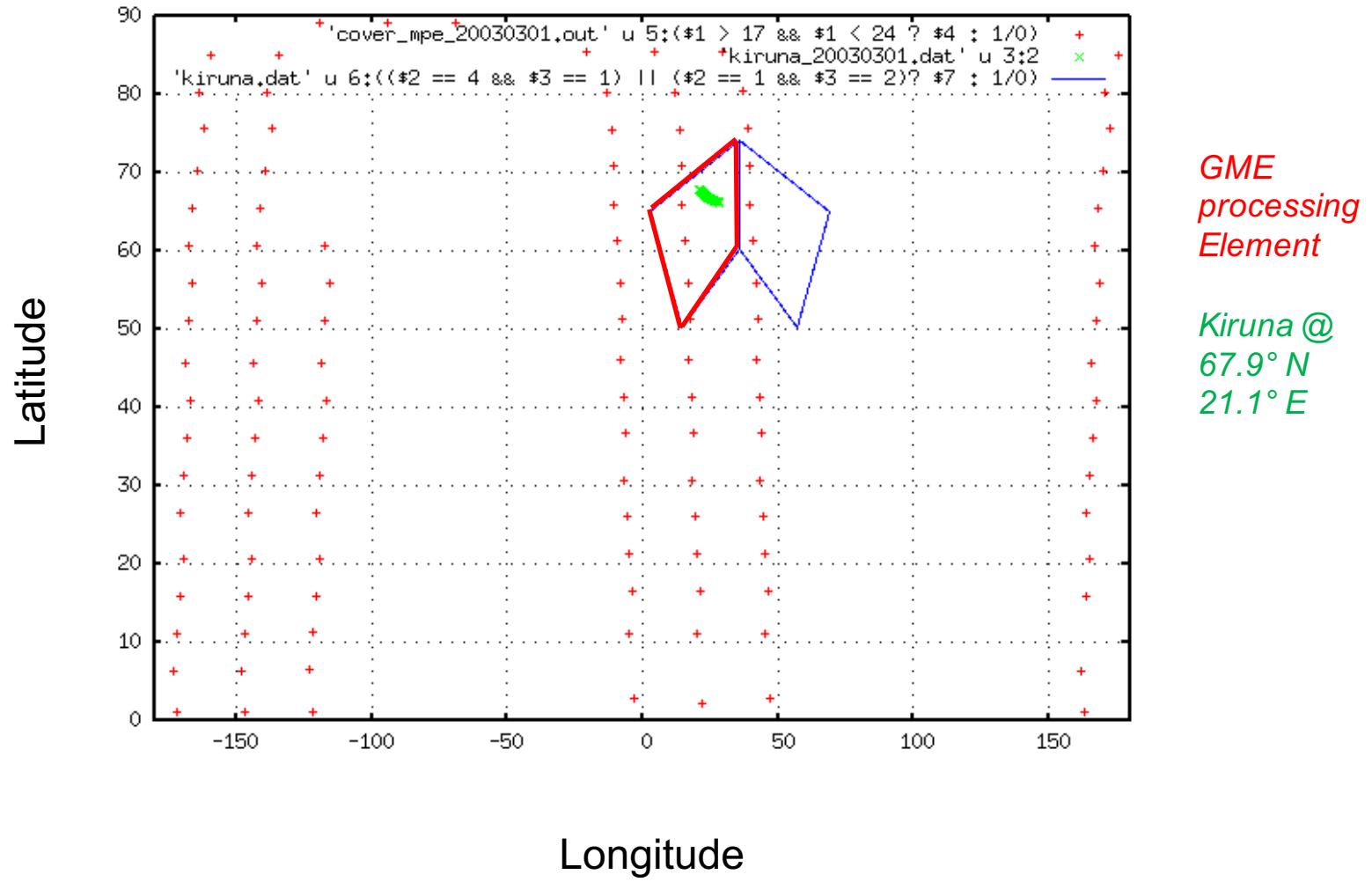
- First-guess winds from GME (=constant)
- MIPAS-SACADA analysis as tracer first-guess
- 6h assimilation window, no background in cost function
- Single processing unit for comparison with Kiruna soundings
- No background in cost function

Dot-product test for twin experiment:

$$\delta x \delta x' \equiv \delta y \delta y$$



MIPAS trace gas observations and Kiruna soundings within 6h window



Detailed analysis of winds by Kiruna soundings

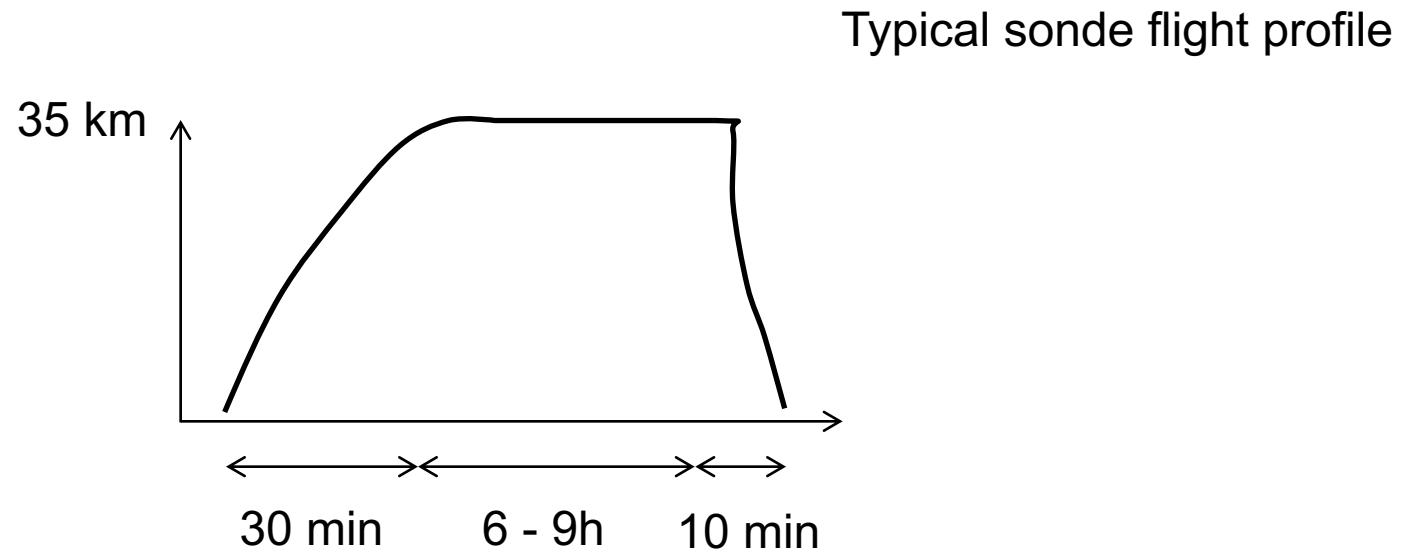
- CNRS Kiruna GPS radiosondes
 - 0 – 30 km altitude level
 - Flight time 6 – 9 h (dismiss drop period)
 - 8 soundings in March 2003

Date	Start time	Stop time	Assim. window	
2003-03-01	18:20	23:59	18-24h	★
2003-03-04	12:53	17:30	12-18h	
2003-03-06	05:54	11:02	06-12h	★
2003-03-20	18:14	09:39 (day after)	18-24h	★
2003-03-23	14:12	18:11	12-18h	
2003-03-25	08:57	11:19	06-12h	★
2003-03-30	15:27	18:50	12-18h	
2003-03-31	23:39	04:38 (day after)	18-24h	★

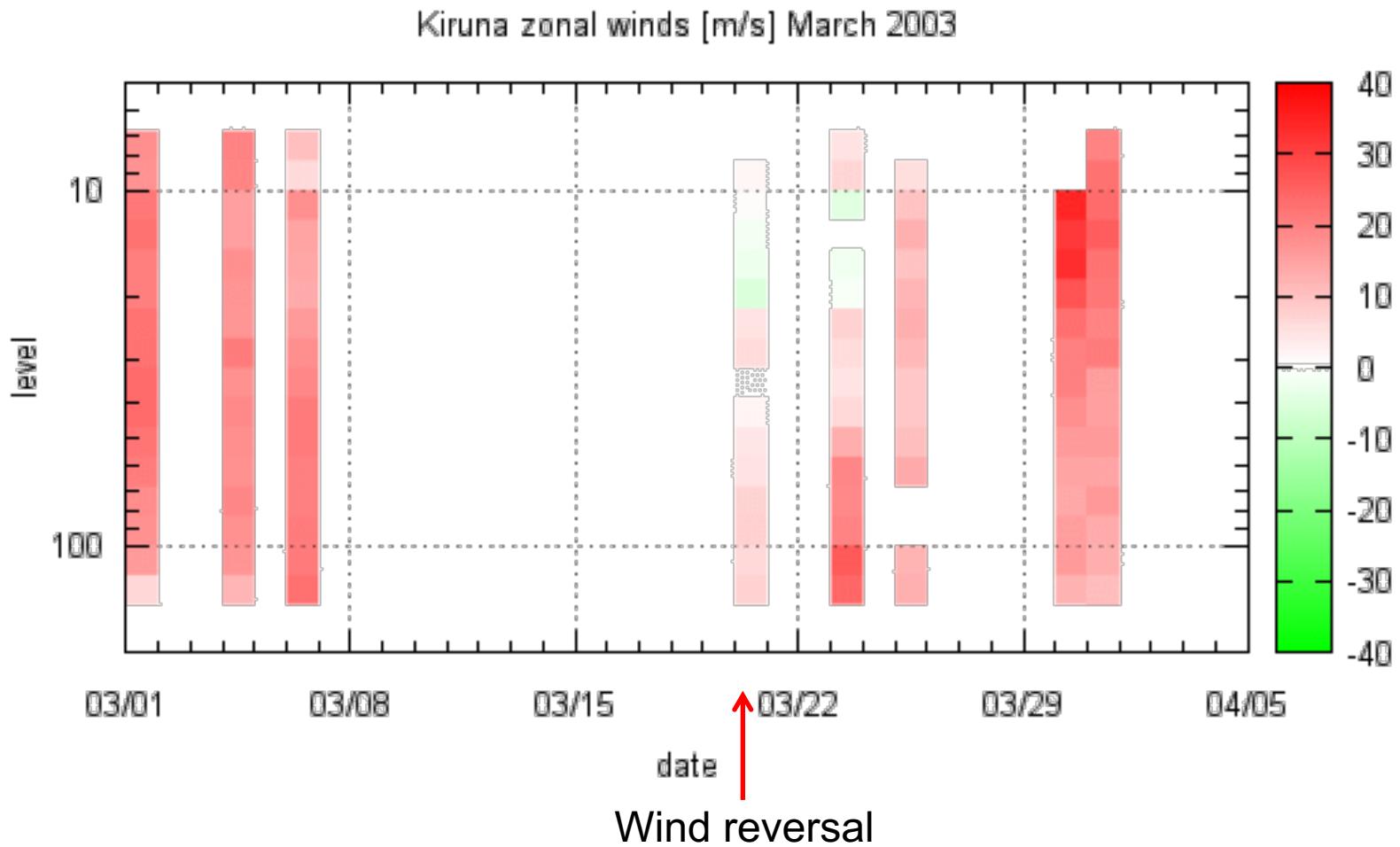
MIPAS
coverage

Detailed analysis of winds by Kiruna soundings

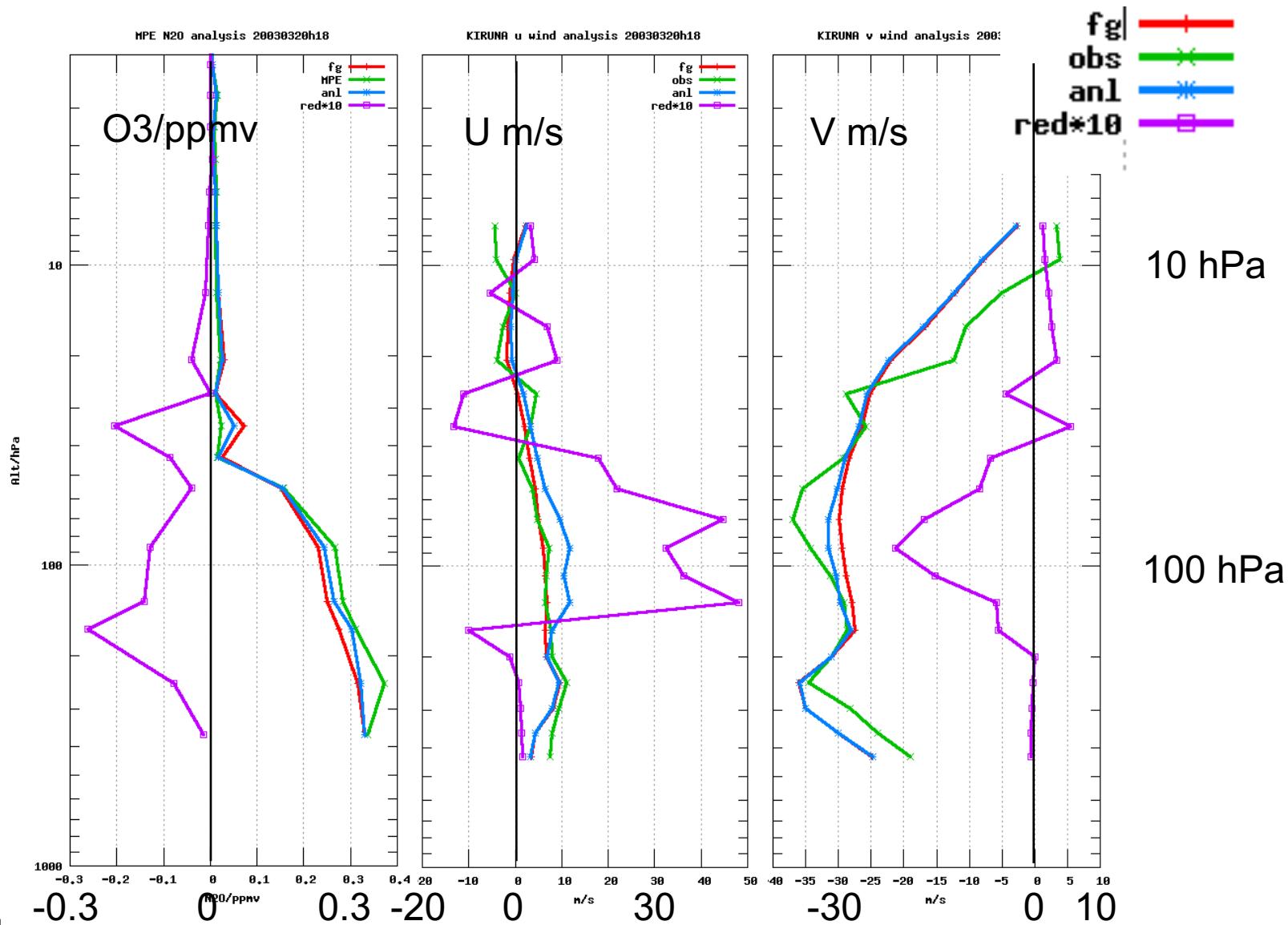
- H-operators
 - Binning of MIPAS w.r.t. model g.p. and time (4D trajectory)
 - Binning of (stationary) Kiruna flight-profile (3D trajectory)



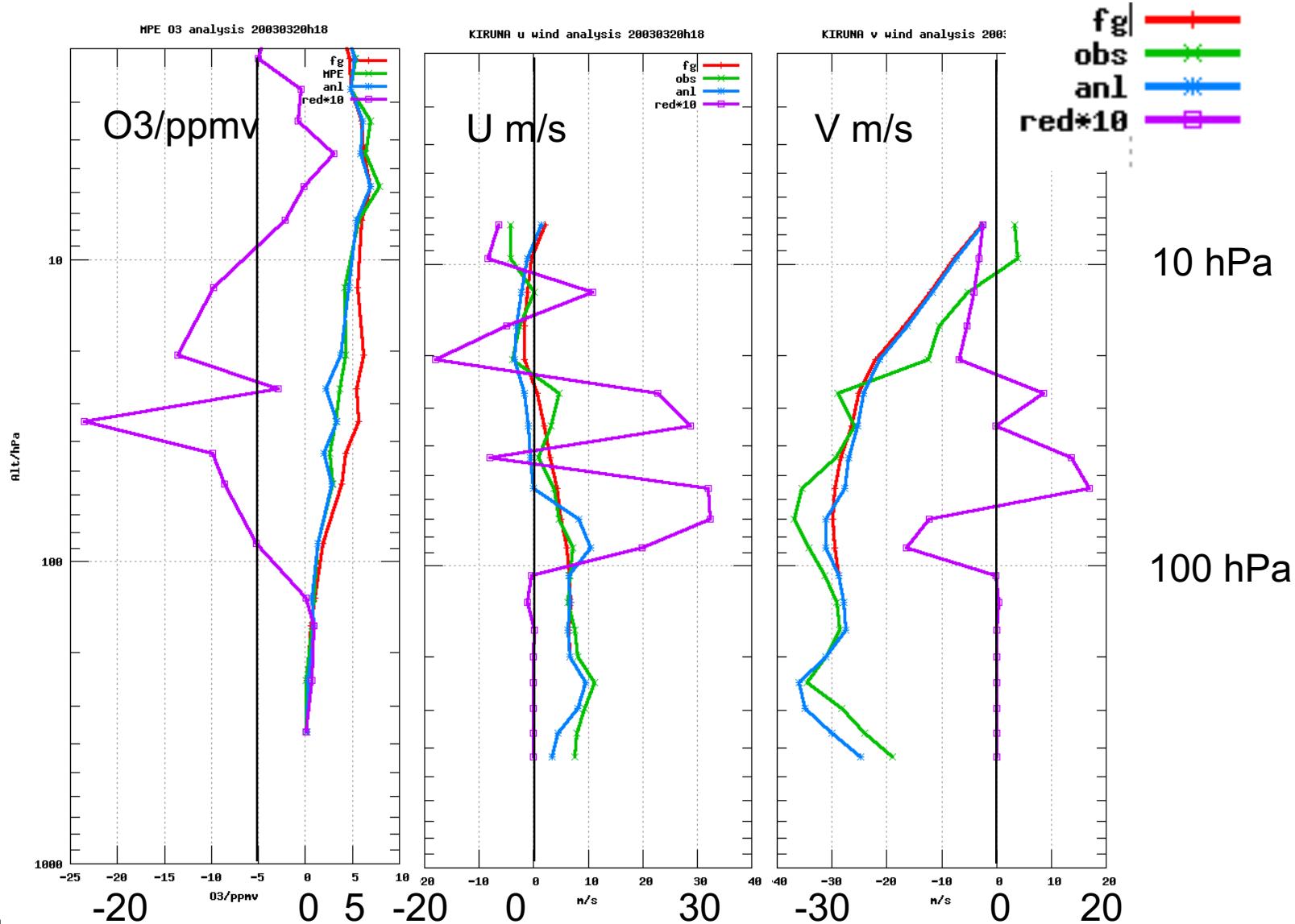
Kiruna soundings March 2003



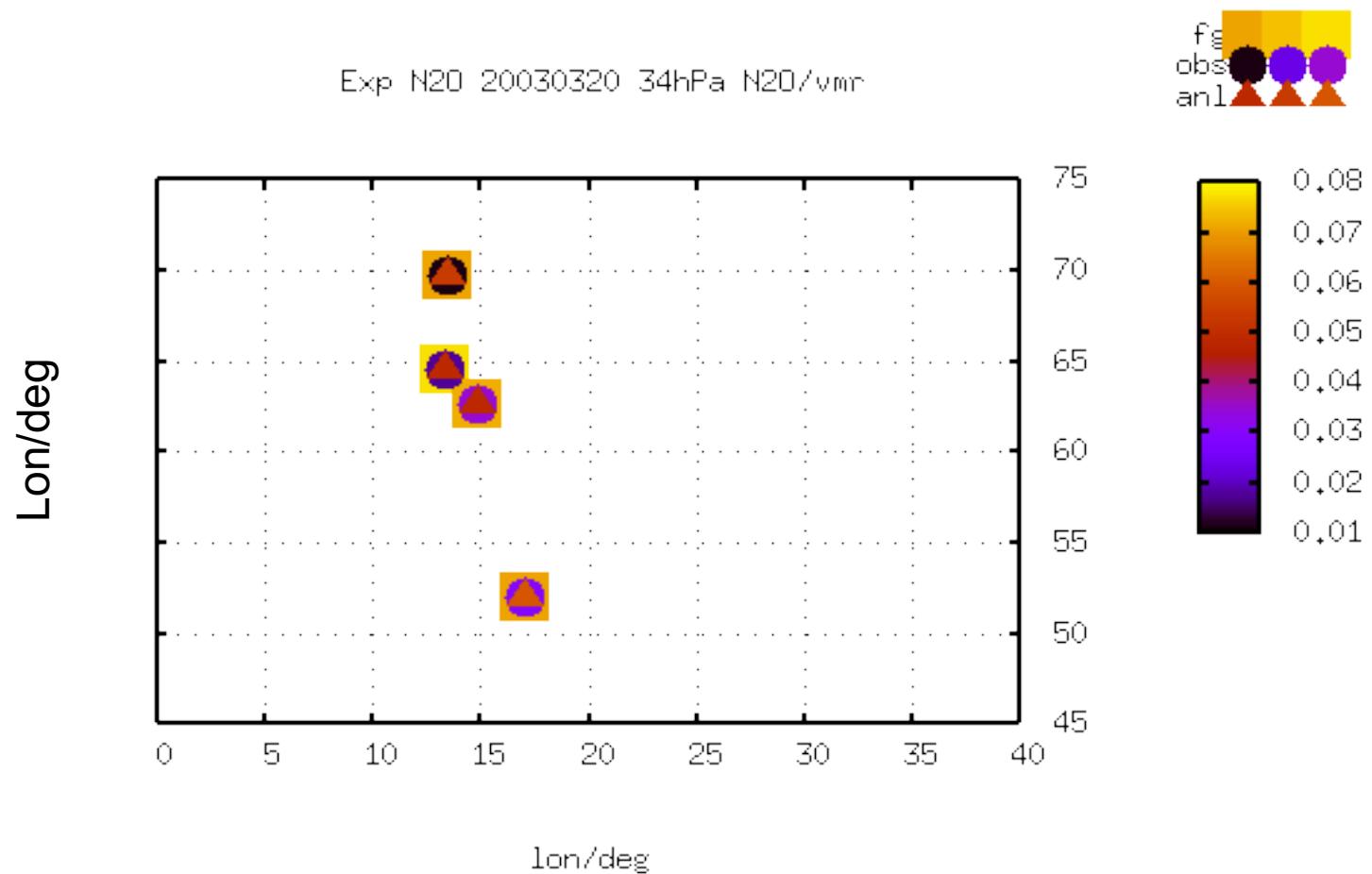
Wind & N₂O tracer results for March 20, 2003



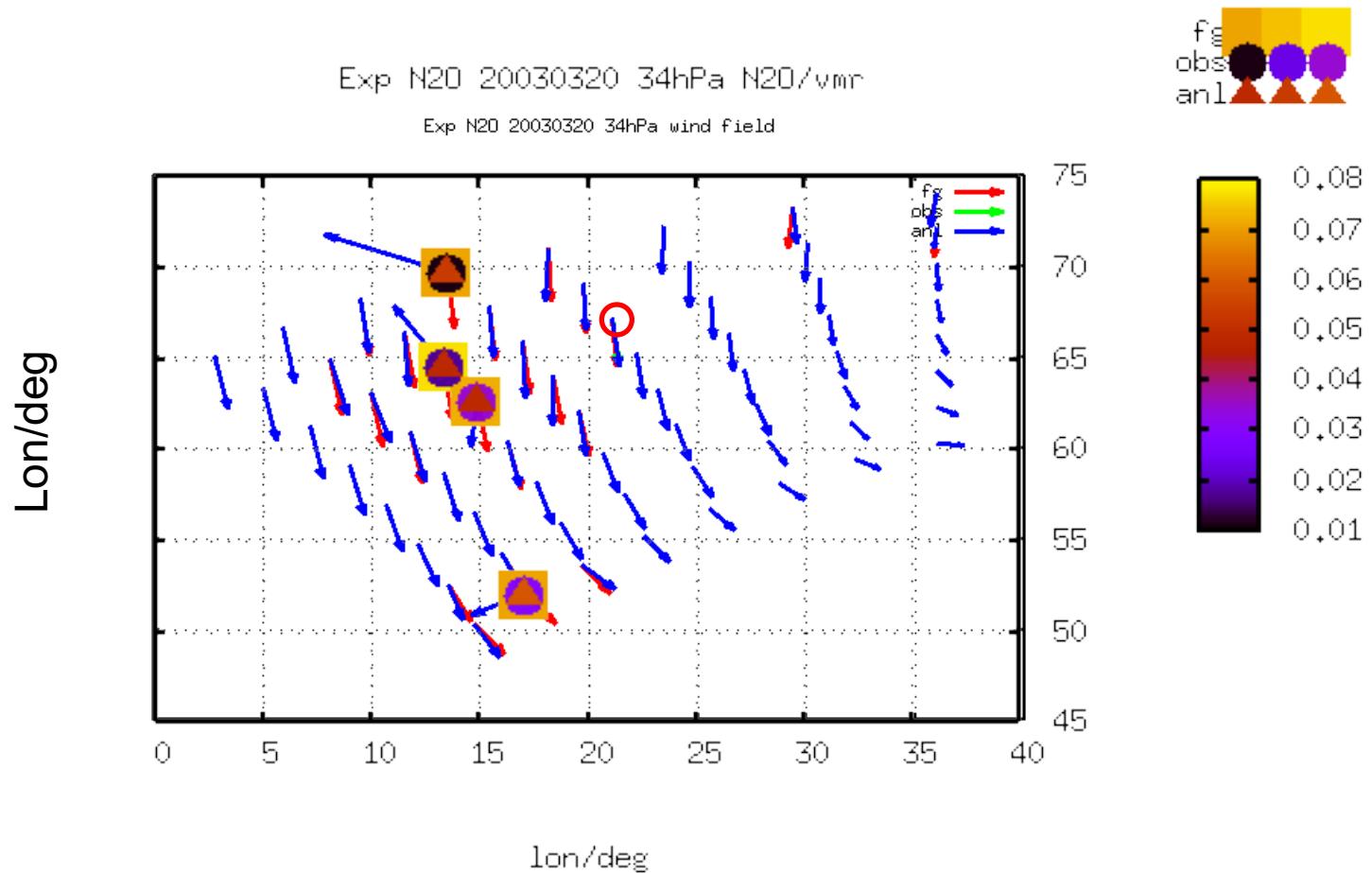
Wind & O3 tracer results for March 20, 2003



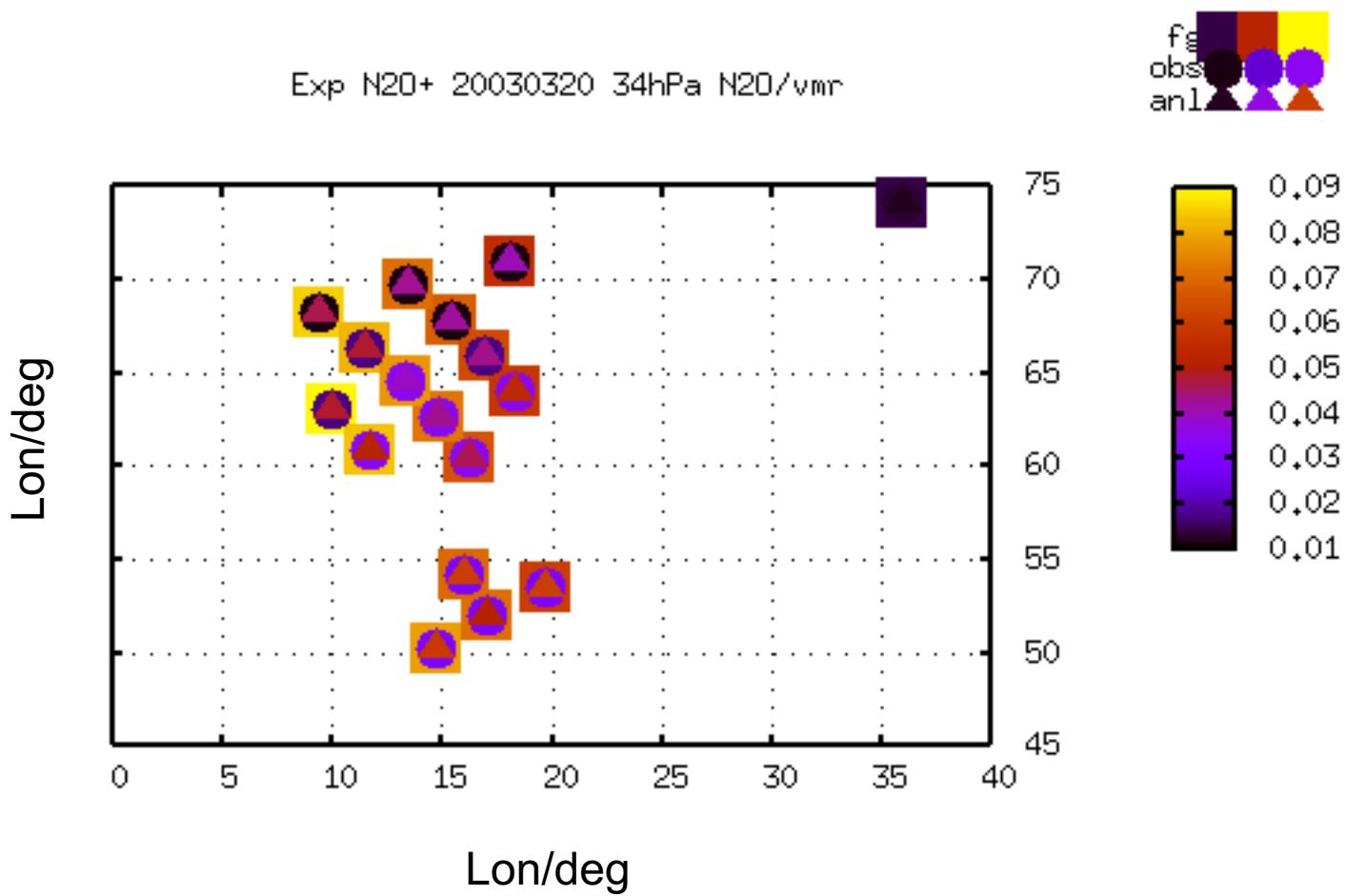
Effect of inversion on wind and tracer fields



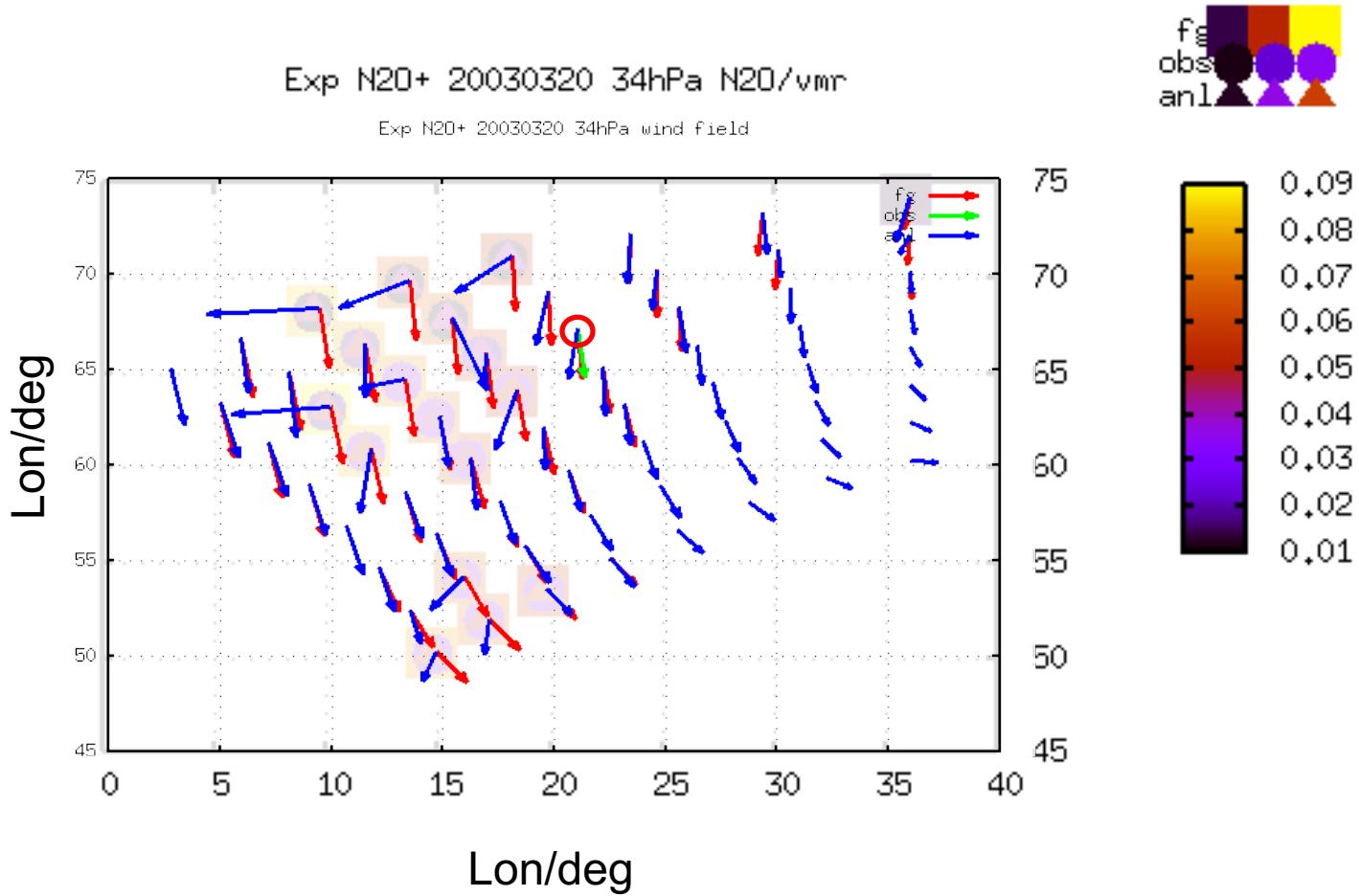
Effect of inversion on wind and tracer fields



MIPAS with extended influence radius



Effect of inversion on wind and tracer fields



Discussion

- Improving wind fields is a challenge:
 - First-guess winds were already close to observations
 - Rapid loss of skill compared to chemical tracers
- Caveats:
 - Missing background-error
 - Constant wind assumption
 - MIPAS 6h coverage



Summary

- Analyzed winds have positive effect on tracer fields
- Comparison to Kiruna shows:
 - More vertical variability (good)
 - Impact on winds strongly depends on H-operator/covariance
- To do:
 - Application to other soundings/latitudes
 - Add linear wind tendency term

