



# **Strengthening of the tropopause inversion layer during the 2009 sudden stratospheric warming in the MERRA-2 analysis**

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# Outline

We use the MERRA-2 analysis and model simulations to study the behavior of the polar tropopause during the 2009 major sudden stratospheric warming (SSW)

- Definitions/conventions
- The tropopause inversion layer during SSWs in MERRA-2
- The mechanisms involved – an analysis of model forecasts
- Summary

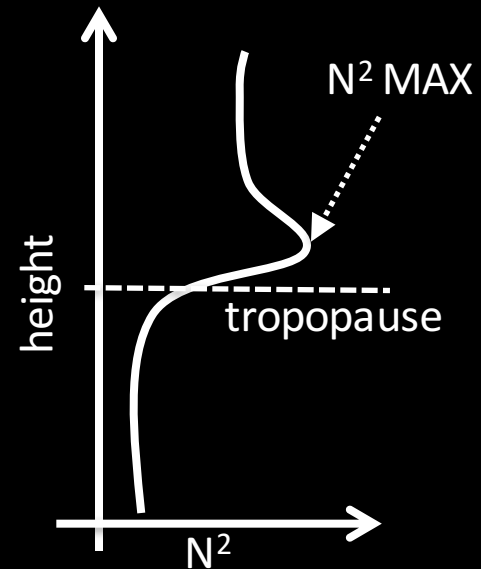
# MERRA-2

(features most relevant to this study)

- Resolution:  $0.635^\circ$  by  $0.5^\circ$  longitude by latitude, 72 layers,  $\sim 1$  km in the vertical near the tropopause
- Observations include hyperspectral IR radiances, GPS radio occultation, radiosondes and aircraft – capable of resolving vertical structures near the tropopause

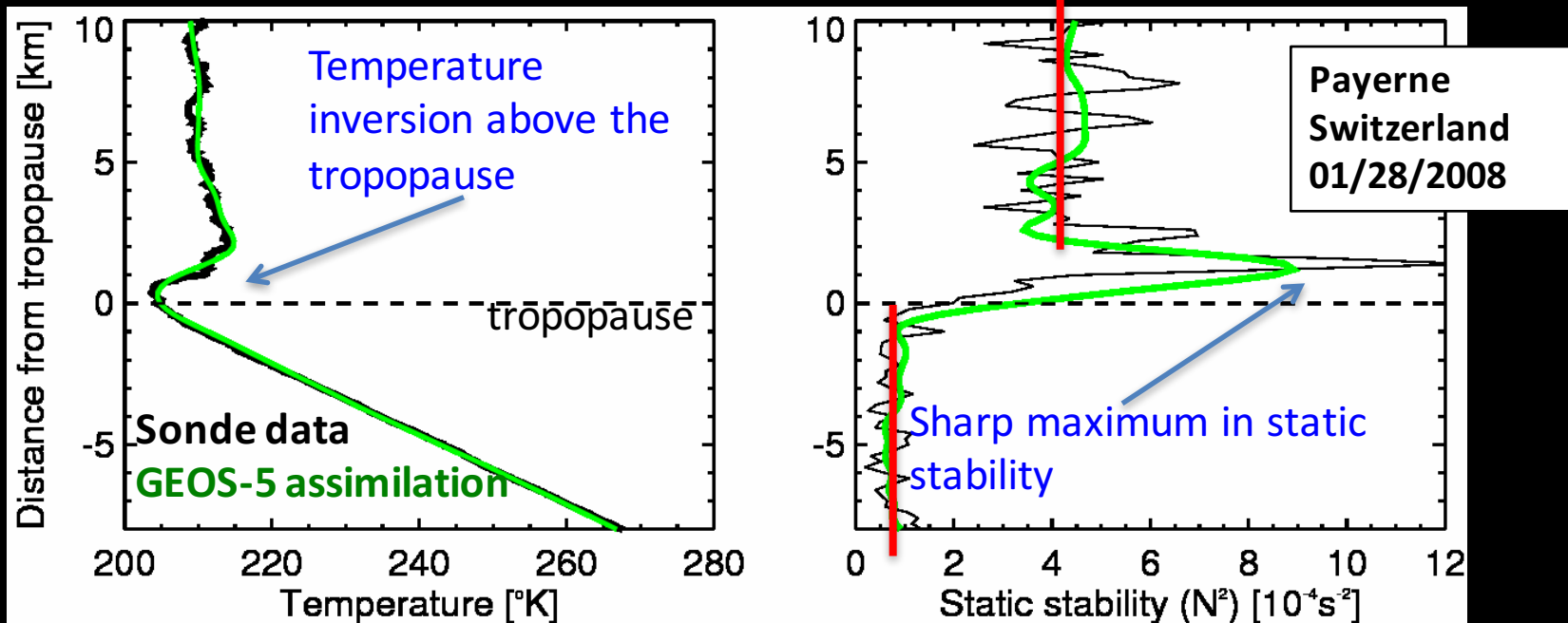
# Definitions/conventions

- Tropopause – standard WMO definition
- Profiles are averaged in tropopause-based coordinates
- Measure of TIL magnitude: Maximum Brunt–Väisälä buoyancy frequency squared,  $N^2$  MAX, within 3 km above the tropopause



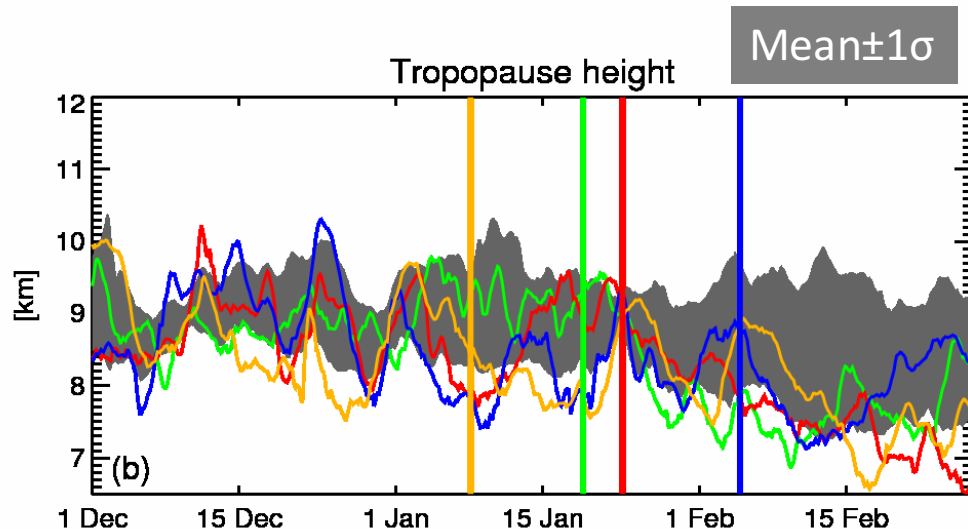
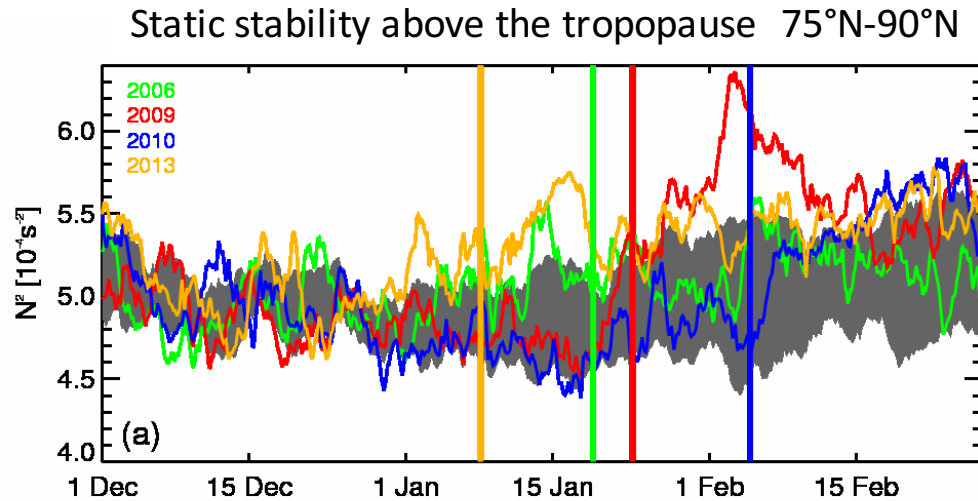
$$N^2 = \frac{g}{\theta} \frac{d\theta}{dz}$$

# What is the Tropopause Inversion Layer (TIL)?



- Positive temperature lapse rate in a 2 – 3 km layer above the tropopause
  - Sharp maximum in static stability
    - Troposphere:  $N^2 \sim 1 \times 10^{-4} s^{-2}$
    - Stratosphere:  $N^2 \sim 4 \times 10^{-4} s^{-2}$
    - TIL:  $N^2 > 5.5 \times 10^{-4} s^{-2}$
  - A ubiquitous feature of the extratropical lower stratosphere
  - Importance: consequences for wave propagation and tracer transport
- In the past, analyses produced weak TIL compared to models

# Motivation: Behavior of static stability during sudden stratospheric warming events



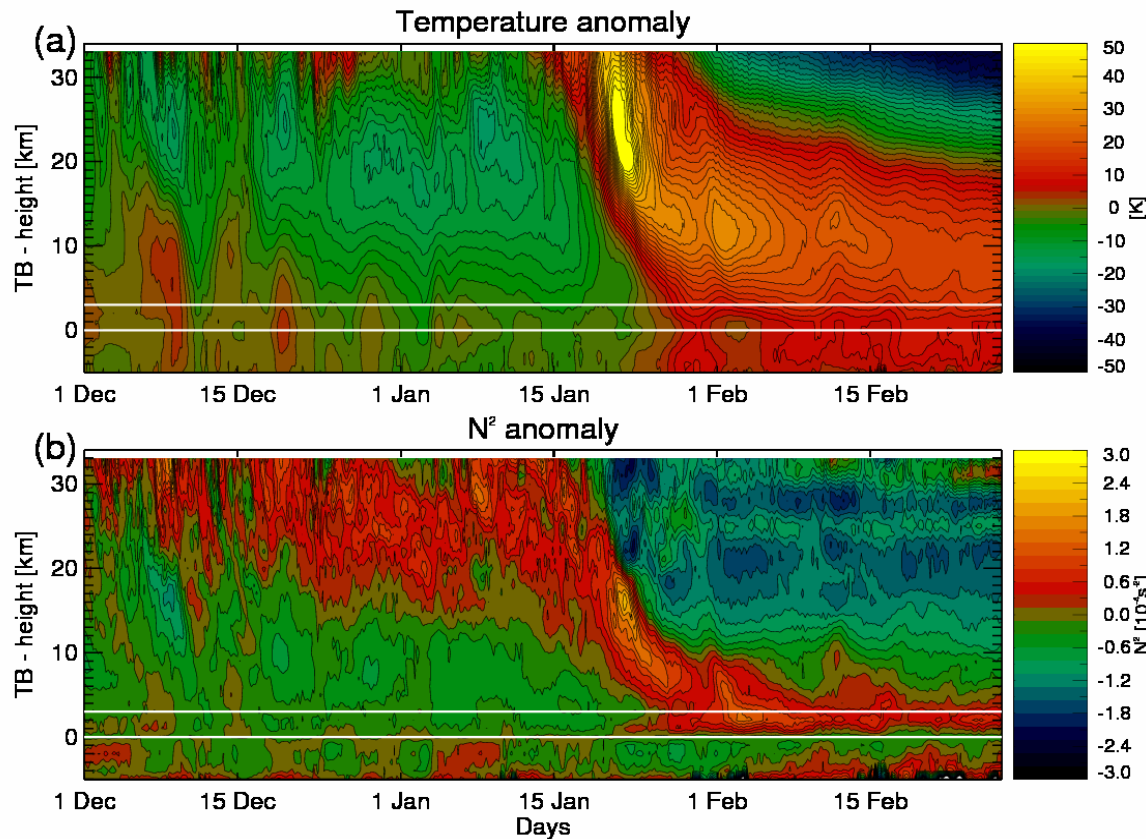
- Static stability above the high-latitude tropopause increases sharply following strong SSWs
- The tropopause height decreases
- Consistent with a result based on GPS-RO data (*Grise et al. 2010*)

The TIL gets stronger during SSWs. What is the mechanism?

We investigate the 2009 case

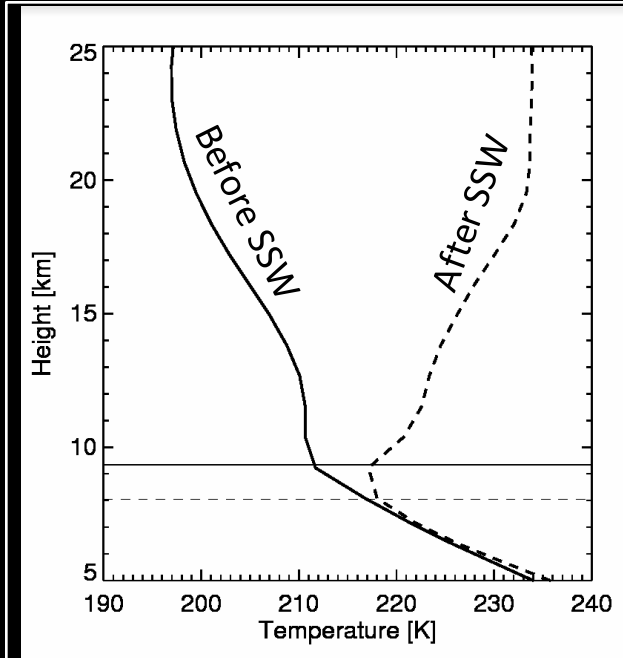
Data: MERRA-2 reanalysis

# High-latitude TTL gets stronger and the tropopause drops during the SSW



75°N-90°N average

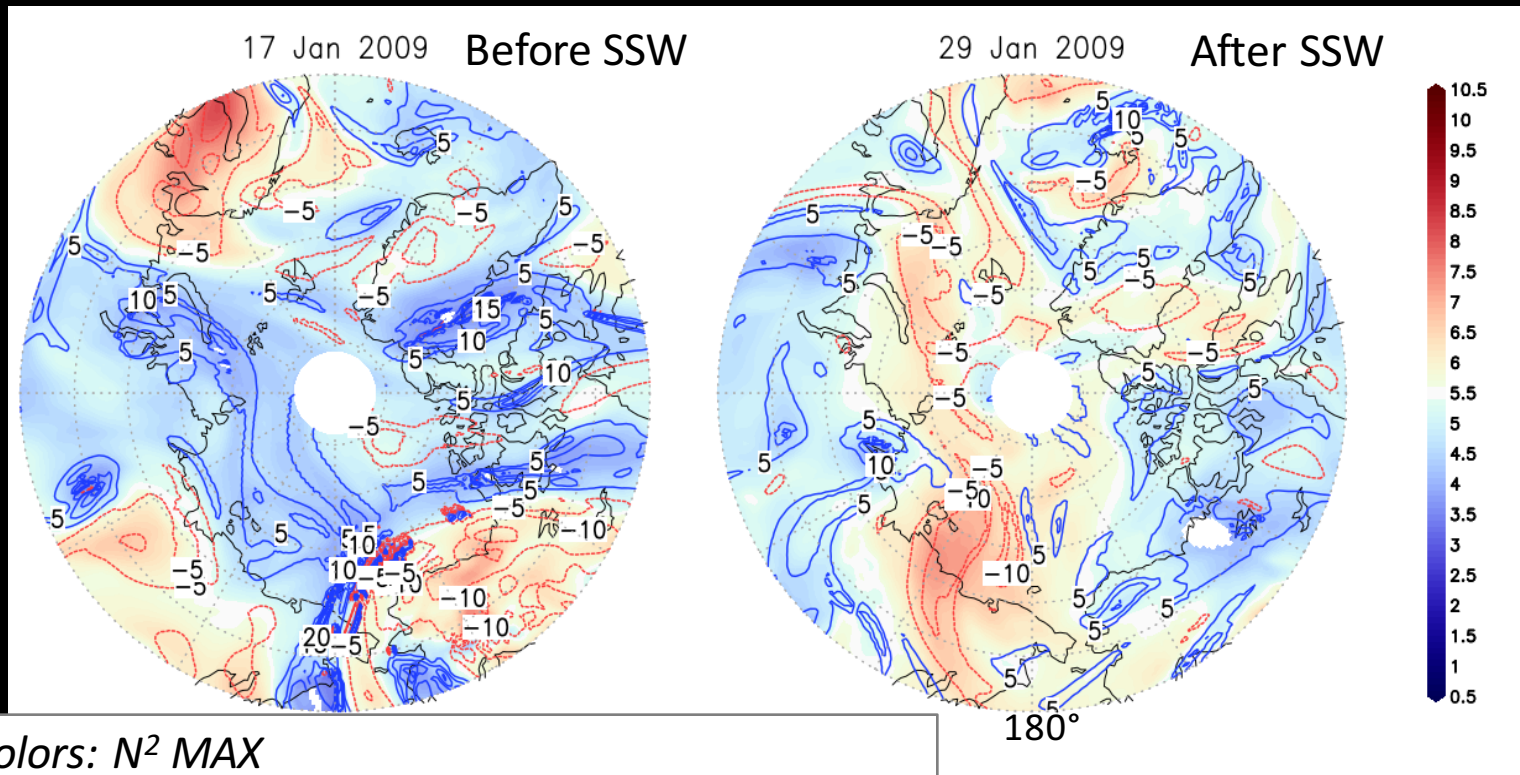
- Rapid downward propagating increase in temperature associated with the SSW
- Positive anomaly in static stability develops along the lower edge of the temperature anomaly



As the lower stratosphere gets warmer the lapse rate decreases → the tropopause moves downward. Most of the apparent warming of the upper troposphere is the result of the decreased tropopause height

# Static stability and circulation at the tropopause

$N^2$  MAX and relative vorticity at the tropopause on 2 days



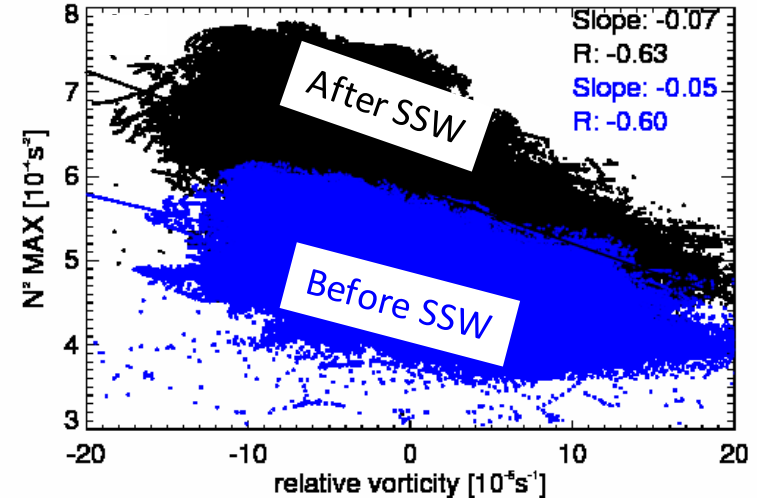
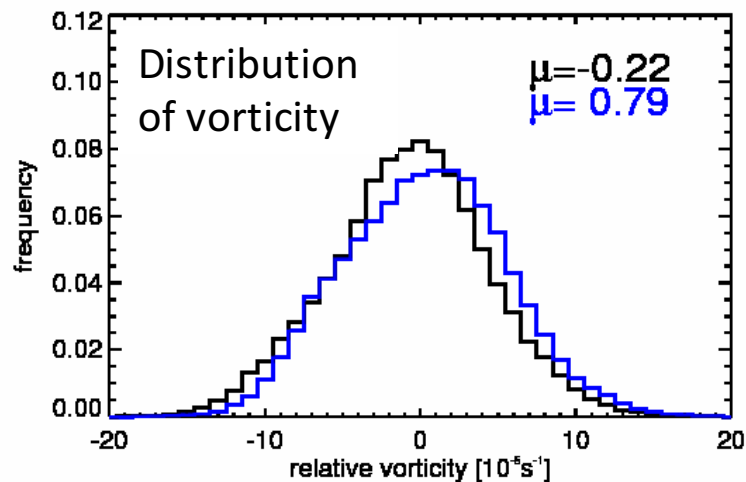
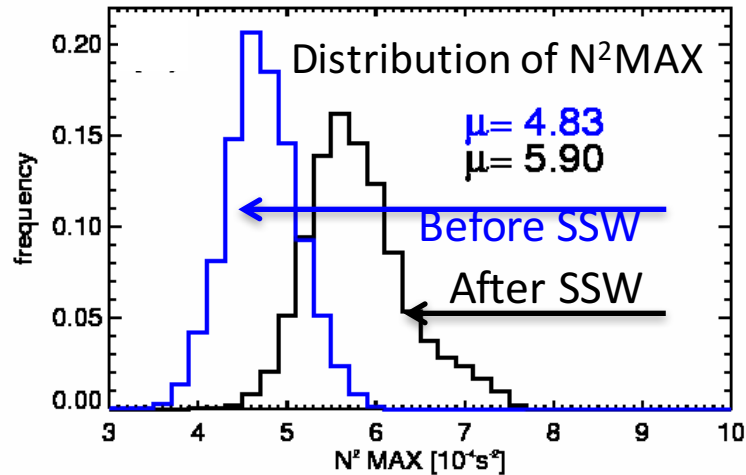
Colors:  $N^2$  MAX

Contours: relative vorticity *cyclonic* *anticyclonic*

- Areas of high/low  $N^2$  MAX coincide with anticyclonic/cyclonic circulation at the tropopause. This is consistent with previous studies
- Again, higher  $N^2$  MAX in high latitudes after the onset of the SSW



# How much of the static stability increase is due to anticyclonic wave breaking (more negative relative vorticity)?



## A back-of-the-envelope calculation

$$[\text{average slope}] \times \Delta[\text{vorticity}] = -0.06 \times 1.01 = -0.0606 \approx 6 \% \text{ of } \Delta N^2 \text{ MAX}$$

At most 6 % of the stability increase is attributed to the decreased vorticity.

75°N-86°N between 1-15 January 2009 (blue) and 28 January - 11 February (black)

**Another mechanism is dominant**

# The budget equation

Vertical component of  
the residual circulation

Diabatic heating rate

$$\partial_t \overline{N}_{MAX}^2 \approx -\overline{N}_{MAX}^2 \partial_z \overline{w^*} |_{MAX} + g \partial_z (\overline{\theta}^{-1} \overline{Q}) |_{MAX}$$

*Birner 2010*

Convergence term

# The budget equation

Vertical component of  
the residual circulation

Diabatic heating rate

$$\partial_t \overline{N}_{MAX}^2 \approx -\overline{N}_{MAX}^2 \partial_z \overline{w^*} |_{MAX} + g \partial_z (\overline{\theta}^{-1} \overline{Q}) |_{MAX}$$

*Birner 2010*

Convergence term

Small

# The budget equation

Vertical component of  
the residual circulation

Diabatic heating rate

$$\partial_t \overline{N}_{MAX}^2 \approx \underbrace{-\overline{N}_{MAX}^2 \partial_z \overline{w^*}}_{\text{Convergence term}} \big|_{MAX} + \underbrace{g \partial_z (\overline{\theta}^{-1} \overline{Q})}_{\text{Small}} \big|_{MAX}$$

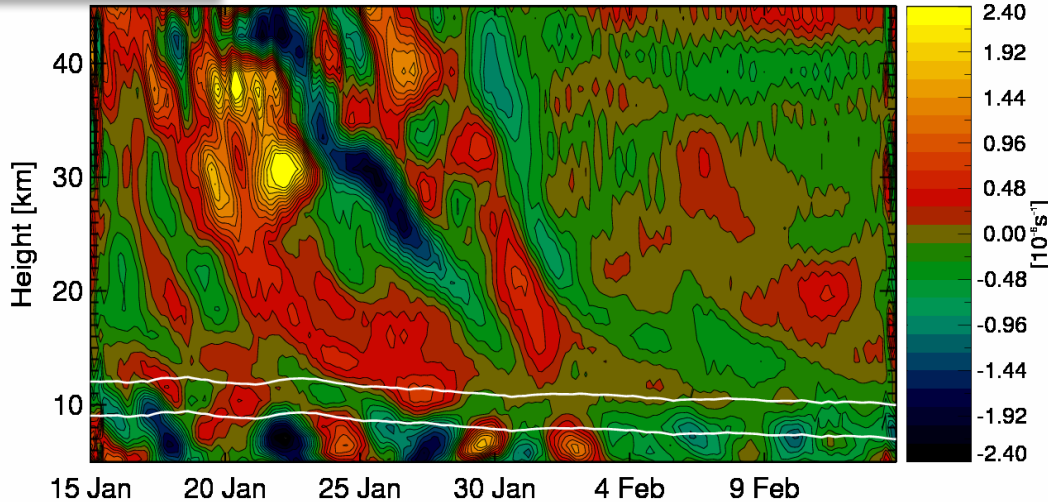
*Birner 2010*

$$-\partial_z \overline{w^*}$$

Convergence term

Small

$w^*$  convergence 75°N-90°N



SSW : downward propagation of

disturbance of the vortex

acceleration of downward motion ( $w^*$ )

Convergence of  $w^*$

**Hypothesis:** the main driver of the increase in static stability (TIL strengthening) during the SSW is the enhanced vertical convergence of the downward residual circulation

$$\partial_t \overline{N}_{MAX}^2 \approx -\overline{N}_{MAX}^2 \partial_z \overline{w}^*|_{MAX}$$

But in MERRA-2 (or any other analysis), by construction

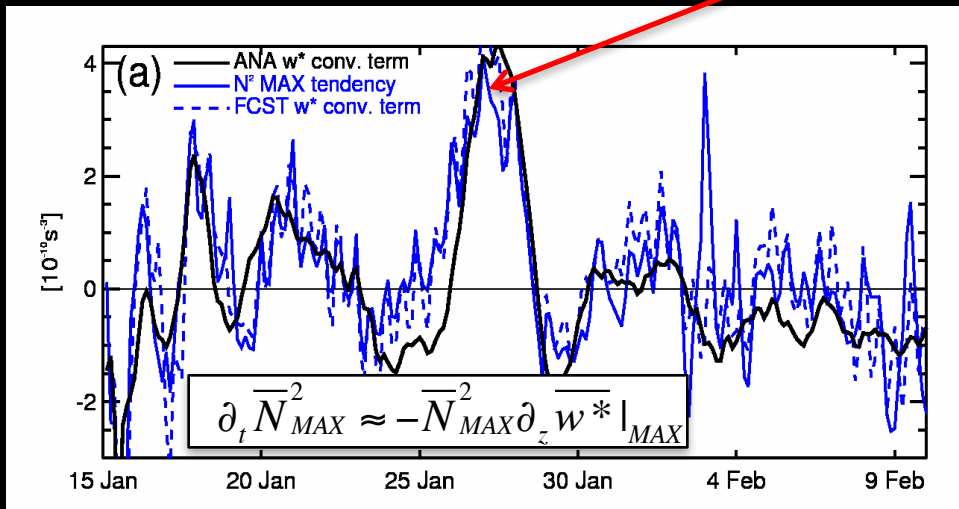
Tendency = tendency due to model physics + tendency due to data insertion

We can't expect the budget to close in the analysis because the left-hand side contains the (unphysical) analysis correction term ('the analysis increment')

Instead, we use an ensemble of forecasts to evaluate the budget:  
23 10-day forecasts initialized from MERRA-2 between 14 January and 4 February 2009

# Computing the $N^2$ MAX budget

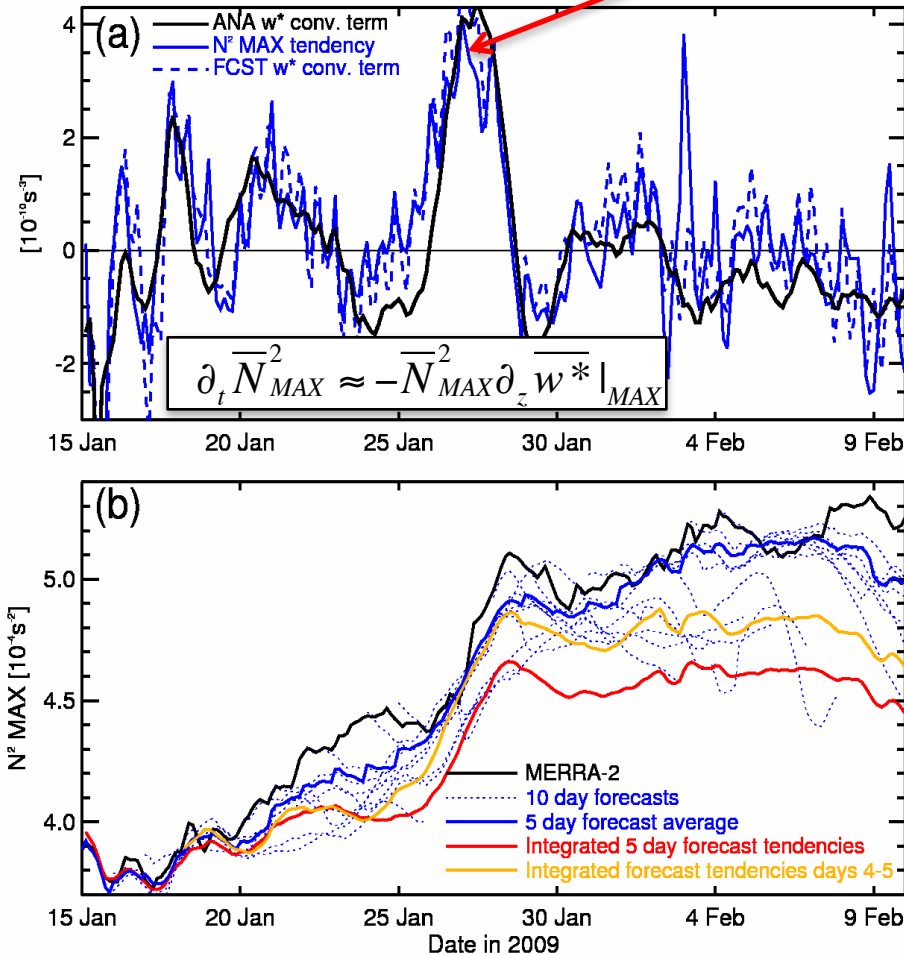
The TIL's rapid strengthening



- The budget closes: Most of the TIL strengthening in the **model** is explained by enhanced vertical convergence of the vertical residual velocity (ensemble mean)
- The convergence in the MERRA-2 analysis agrees with that in the model

# Computing the N<sup>2</sup> MAX budget

The TIL's rapid strengthening

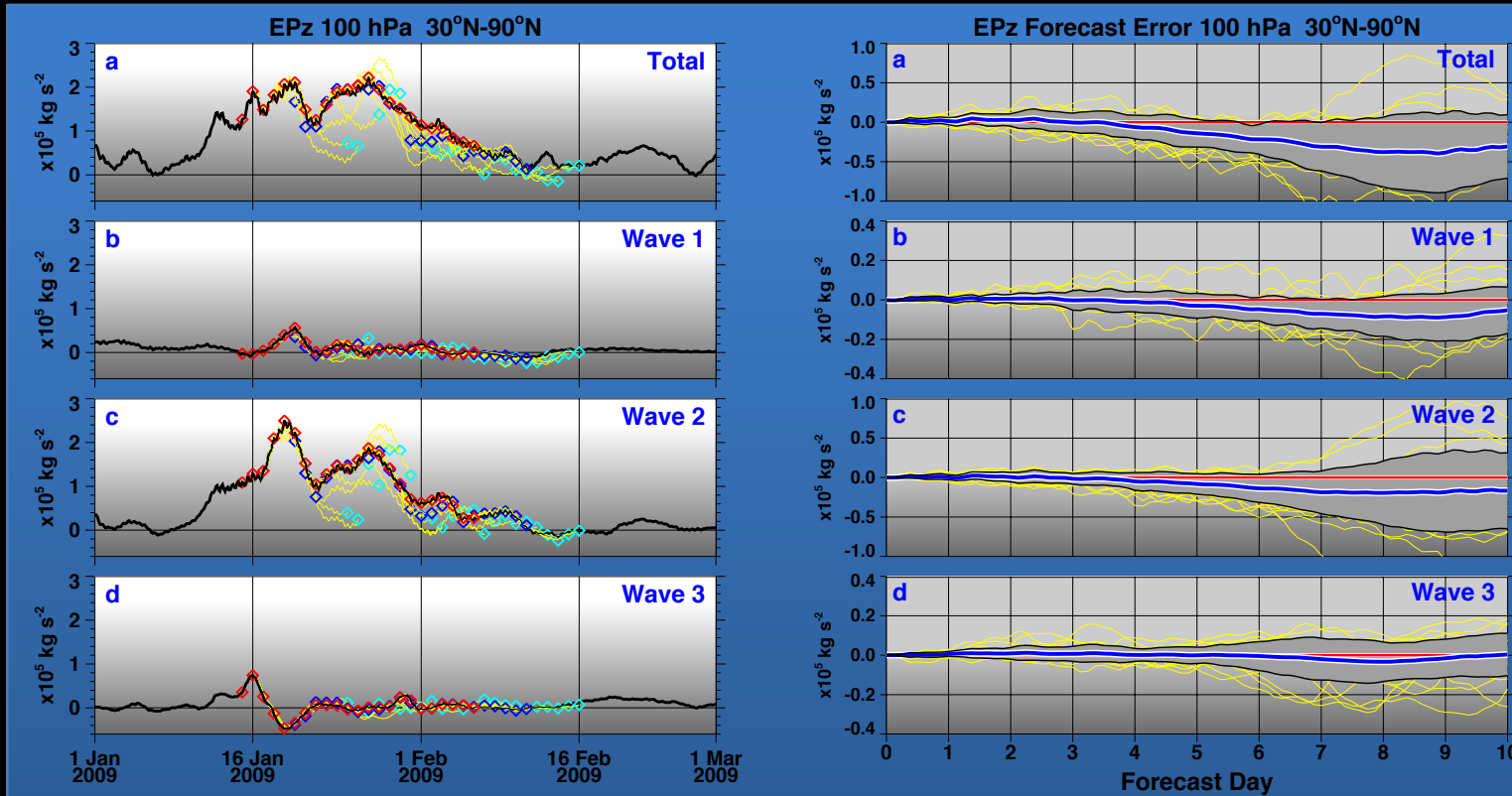


- The budget closes: Most of the TIL strengthening in the **model** is explained by enhanced vertical convergence of the vertical residual velocity (ensemble mean)
- The convergence in the MERRA-2 analysis agrees with that in the model

**Model simulations** underestimate the strengthening of the TIL compared to MERRA-2: the model tendencies are biased low.

# Why does the model underestimate the static stability increase?

SSWs are forced by planetary wave pulses from the troposphere



100 hPa vertical EP flux from *MERRA-2*, 5-day, 10-day forecasts

Forecast error with respect to analysis

- In 2009 the forcing is dominated by wave 2 (*Harada et al., 2009*)
- There is a tendency for the forecasts to underestimate the wave forcing
- Overall, 5-day forecasts agree with the analysis but there is some underestimation near the wave 2 maximum



# Summary of the results

- MERRA-2 shows strengthening of the polar tropopause inversion layer during major SSW events in agreement with previous studies
- Model simulations reveal that the primary mechanism (in 2009) involves an enhanced convergence of the vertical residual wind at the tropopause
- The model underestimates the TIL's strengthening and wave forcing

# Final remarks

- DA is useful for studies of dynamical coupling in the stratosphere but one has to keep in mind that model tendencies  $\neq$  analysis tendencies: be careful with budget calculations!
- Older data assimilation systems underestimated the TIL compared to models. Now the model produces a weaker TIL
- GMAO plans to double the model vertical resolution in near future. How will the results change? (The optimal vertical-to-horizontal resolution ratio is  $\sim 300$  m/degree *Erler and Wirth 2007*)